



# Temperature regime of drained and natural peatlands in arid and water-logged years

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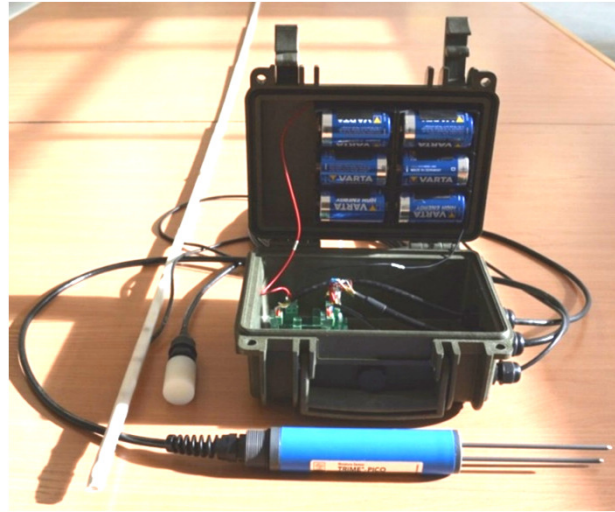
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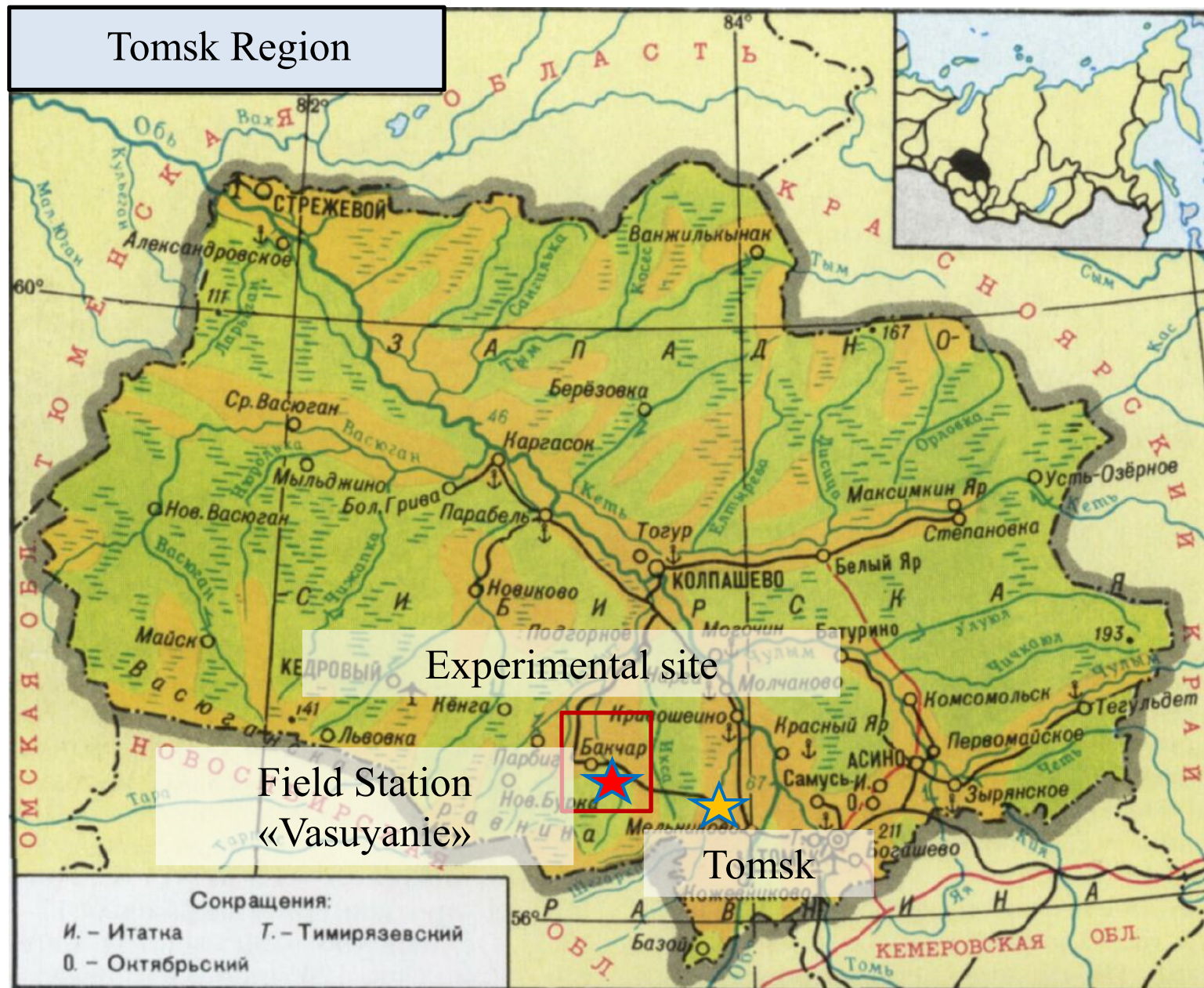
## Atmospheric-soil measuring system (ASMS) *IMCES SB RAS*



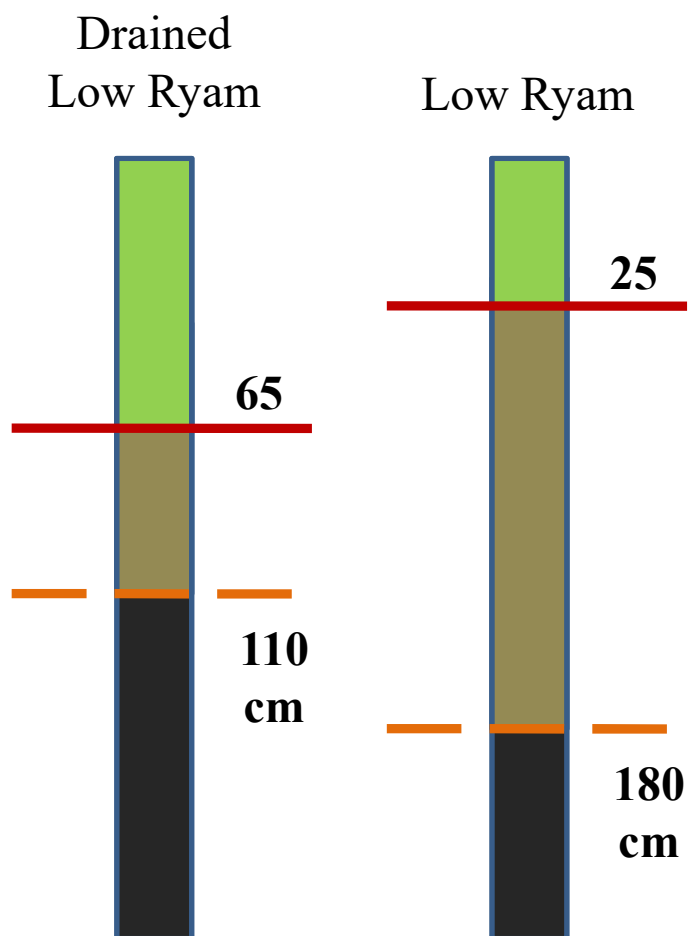
№	Component part	Measuring range	Accuracy
1	Temperature sensor DS18B20	from -55 to +55°C	±0,1°C
2	Air humidity sensor HIH-5031	from 0 to 100%	±3,5%
3	Soil moisture sensor TRIME-PICO32	from 0 to 100%	±2%
4	Atmospheric pressure sensor MPL3115A2	from 500 to 1100 hPa	±4 hPa
5	Wind speed and direction sensor Davis Anemometer 6410	from 0,9 to 78 m/s from 0 to 360°	±5% ±7%
6	Liquid precipitation sensor Davis 7852M	from 0 to 1000 mm/h	±5%
7	Pyranometer 2 channel (albedometer) Range from 0,35 to 9 μm	from 0 to 2000 W/m <sup>2</sup>	±10%
8	The level of groundwater sensor	from 0 to 10,5 m	±1%
9	Snow cover depth sensor	from 0 to 2 m	±0,05 m
10	Ultrasonic sensor of snow cover depth MB7384	from 0 to 4 m	±0,015 m



## Tomsk Region



## Experimental sites



— the peat depth, cm

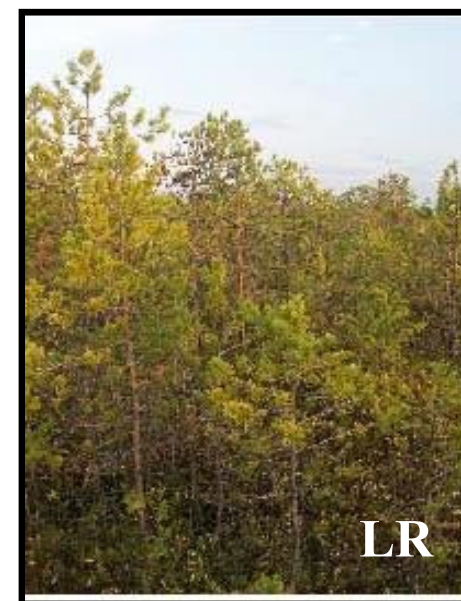
— the average bog water table level, cm

— Anaerobic peat

— Aerobic peat

— Mineral soil

Pine-shrub sphagnum ecosystem



**Measurement period:** 11.09.2011 – 13.10.2018

**Time step:** 15 min

**Temperature accuracy:** 0,1°C

**Depth:** 0 – 240 cm



## The Ped drought index

$$S_i = \frac{\Delta t_i}{\sigma_t} - \frac{\Delta r_i}{\sigma_r}$$

$\Delta t_i = t_i - t_{norm}$  Anomaly of air temperature  
for  $i$ -th time step

$\sigma_t$  Standard deviation  
for air temperature

$\Delta r_i = r_i - r_{norm}$  Anomaly of precipitations  
for  $i$ -th time step

$\sigma_r$  Standard deviation  
for atmospheric precipitation

The greatest  
drought  
in 2012

The greatest  
moistening  
in 2018

Year	Months						
	April	May	June	July	August	September	October
2011	2,60	0,64	1,75	-3,20	-3,49	1,83	1,34
2012	1,80	0,44	4,07	3,42	-1,10	1,56	0,85
2013	0,39	-4,31	-1,30	1,33	0,23	-0,89	0,77
2014	0,57	-1,26	0,54	1,30	1,71	-2,25	-2,53
2015	-1,66	0,16	1,11	0,86	0,16	-2,06	-2,04
2016	0,38	0,55	1,54	-1,54	1,85	3,58	-1,64
2017	0,32	-0,33	-2,14	-0,29	1,33	-0,41	0,54
2018	-2,51	-4,05	-1,74	-1,63	0,56	1,64	1,01

Moistening period



weak ( $1 \leq |S| < 2$ )

moderate ( $2 \leq |S| < 3$ )

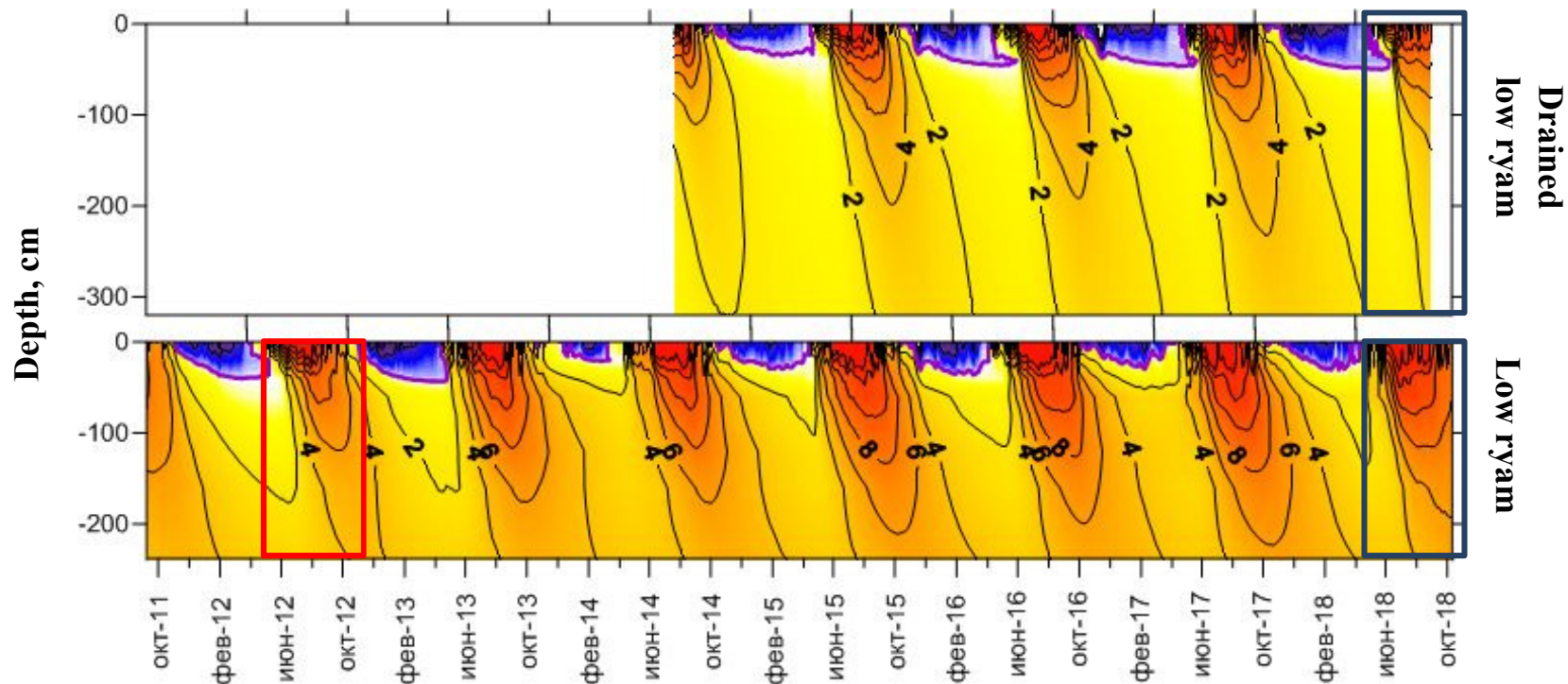
strong ( $3 \leq |S| < 4$ )

Extreme events ( $|S| \geq 4$ )

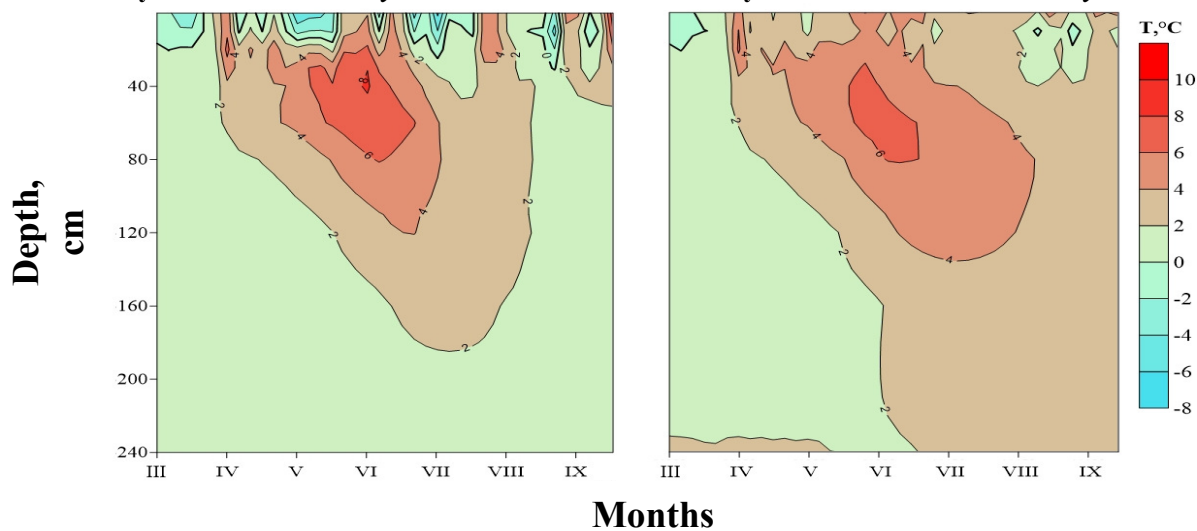
Drought period



# Peat soil temperature profile 2011-2018



$$\Delta T = T_{\text{low ryam 2018}} - T_{\text{low ryam 2012}} \quad \Delta T = T_{\text{low ryam 2018}} - T_{\text{drained low ryam 2018}}$$



# Soil heat flux estimation from observations

Heat-mass transfer equation

$$C \frac{\partial T}{\partial t} = S = -\frac{\partial G}{\partial z} + \Delta F$$

Net change in heat storages

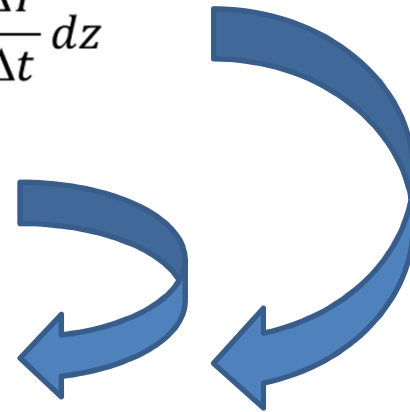
$$\Delta S = \int_{z_0}^{z_1} C \frac{\Delta T}{\Delta t} dz$$

Diffusion heat flux

$$G = -k \frac{dT}{dz}$$

Latent (and other) heat flux

$$F = S - G$$



$$\frac{\partial T}{\partial z} = \frac{T_i^k - T_{i-1}^k}{z_i - z_{i-1}}, \quad \frac{\partial T}{\partial t} = \frac{T_i^k - T_i^{k-1}}{\Delta t},$$

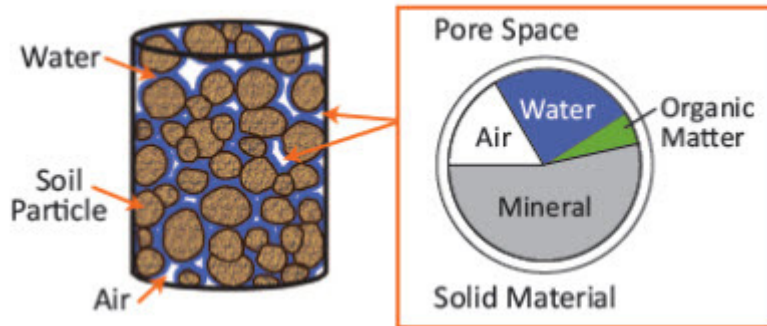


# Soil properties ( A + W + I + S)

$$k = k_a A + k_w W + k_i I + k_s (1 - A - W - I)$$

$$C = c_A \rho_A A + c_w \rho_w W + c_I \rho_I I + c_{s1} \rho_{s1} (1 - \Pi) S1 + c_{s2} \rho_{s2} (1 - \Pi) (1 - S1)$$

Composition of an Unsaturated Soil Sample



$$\rho_A = 1,247 \text{ kg/m}^3$$

$$\rho_w = 1000 \text{ kg/m}^3$$

$$\rho_I = 918 \text{ kg/m}^3$$

$$\rho_{s1} = 1510 \text{ kg/m}^3$$

$$\rho_{s2} = 2700 \text{ kg/m}^3$$

$$c_A = 1005 \text{ J/kg/K}$$

$$c_w = 4182 \text{ J/kg/K}$$

$$c_I = 2150 \text{ J/kg/K}$$

$$c_{s1} = 1880 \text{ J/kg/K}$$

$$c_{s2} = 750 \text{ J/kg/K}$$

$$k_s = 0.05 \text{ W m}^{-1} \text{ K}^{-1}$$

$$k_A = 0.025 \text{ W m}^{-1} \text{ K}^{-1}$$

$$k_w = 0.59 \text{ W m}^{-1} \text{ K}^{-1}$$

$$k_I = 2.25 \text{ W m}^{-1} \text{ K}^{-1}$$

$$\text{at } z \leq WTL \quad W = \Pi, A = 0$$

$$\text{at } z > WTL \quad W = W_0 + (\Pi - W_0)(z/WTL)^2 \quad A = \Pi - W,$$

$$\text{If } T \leq 0, \text{ then } I = W, W = 0, A = \Pi - W.$$

*WTL* – water table level, cm

*W* – water content

*I* – ice content

*A* – air content

*S1* – peat content

*S2* – clay content

*Π* – porosity

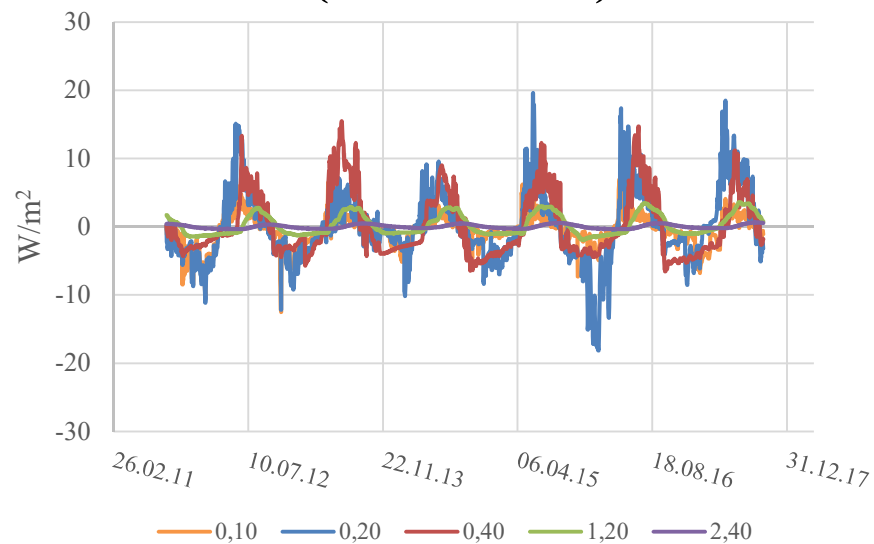
*C* – heat capacity

*ρ* – density

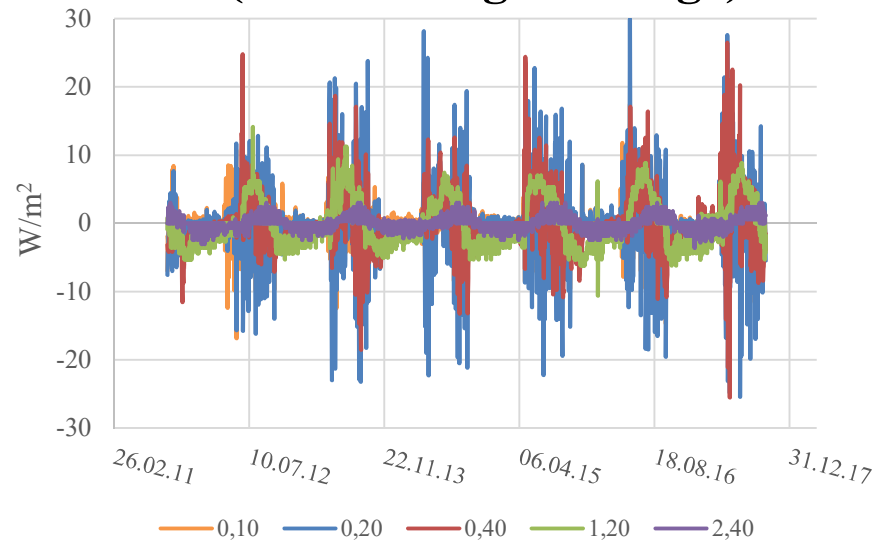
*k* – thermal conductivity

## Estimation results

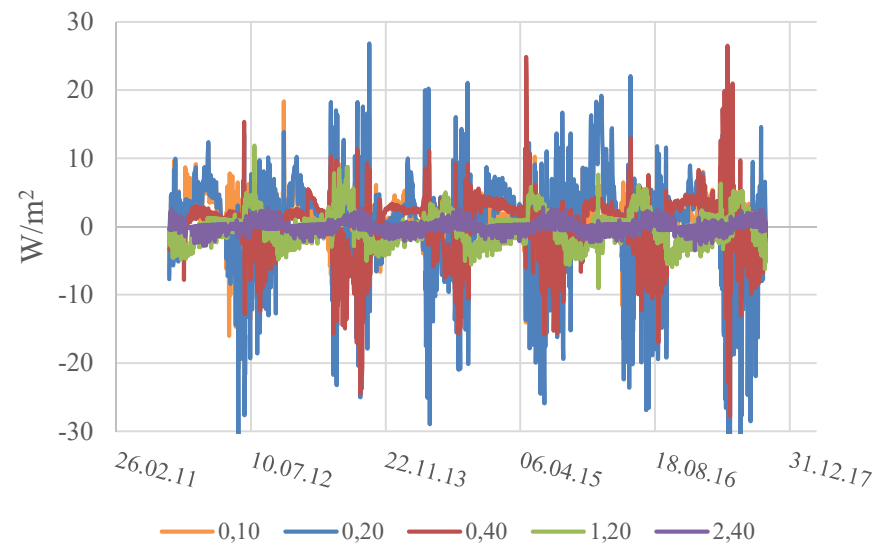
### G (Conduction)



### S (Heat storages change)

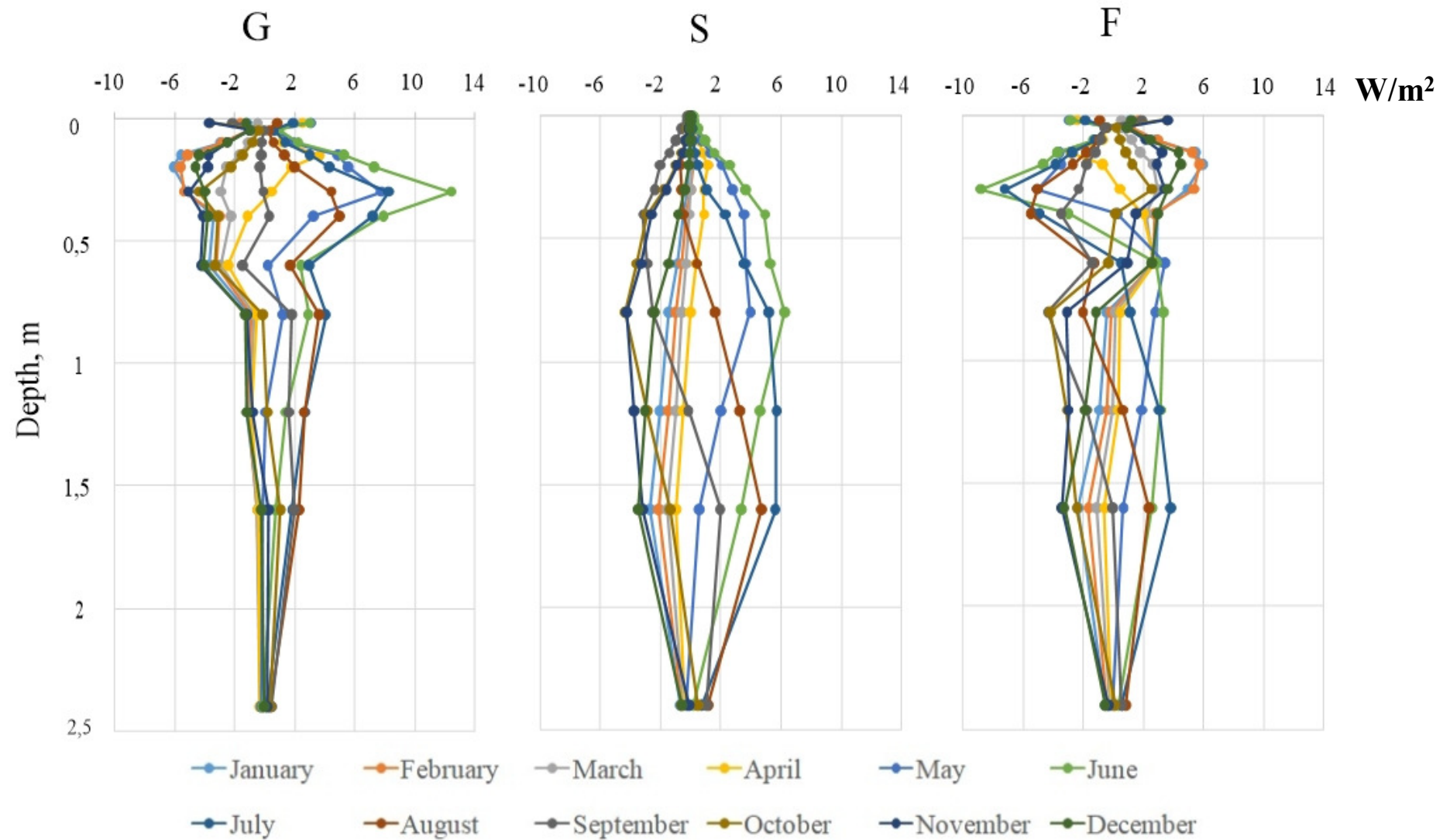


### $F = S - G$ (Latent and nonconductive)

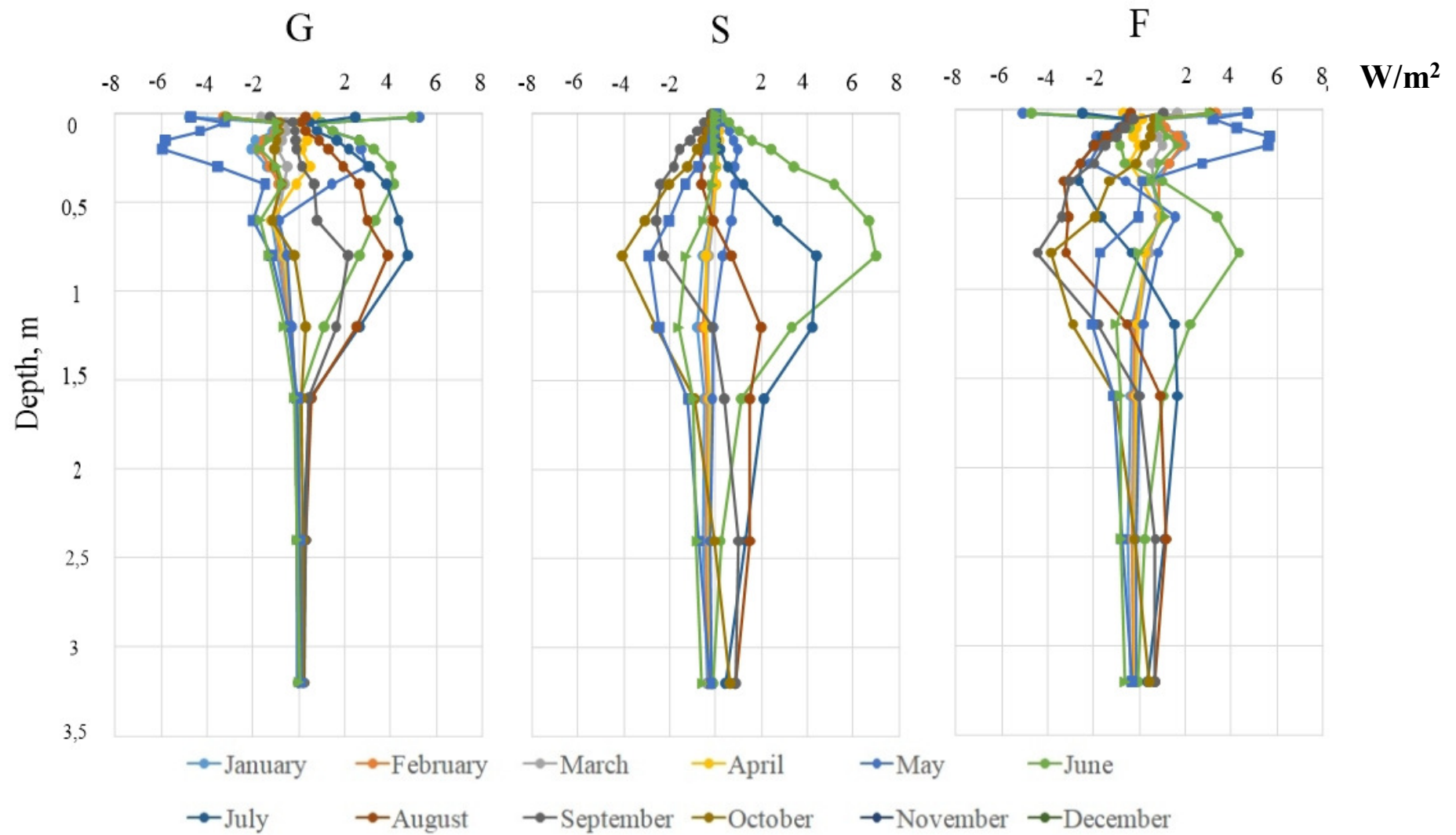




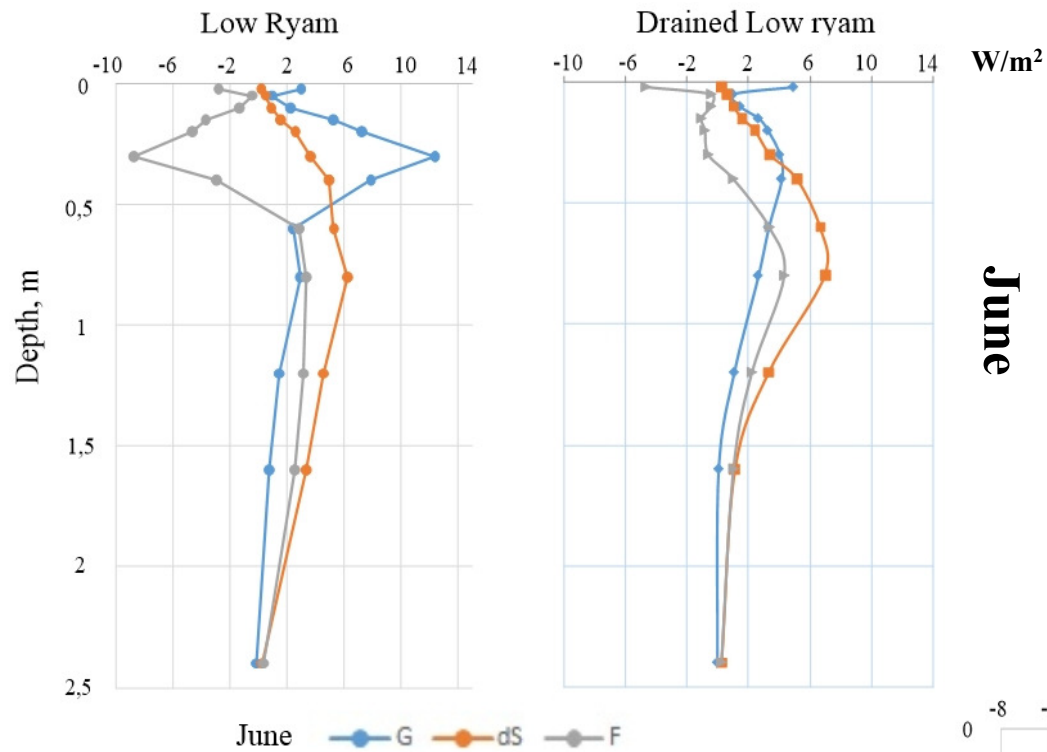
## Soil heat fluxes in low ryam



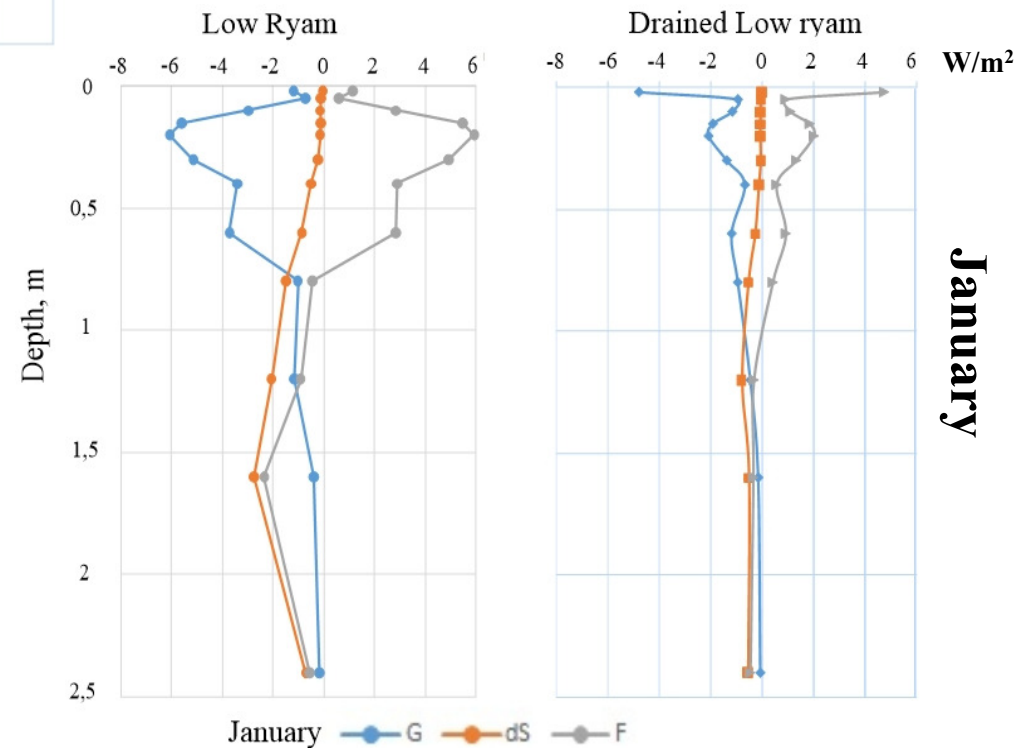
## Soil heat fluxes in drained low ryam







**Monthly averaged heat fluxes for low ryam and drained low ryam (June and January)**



# Conclusions

The analysis of the temperature regime of natural and drained oligotrophic bogs in years with significantly different hydrothermal conditions showed that:

- The main factor controlling the temperature regime of peat deposits is the level of bog waters and, as a result, the peat moisture.

- In an extremely dry and hot year, the temperature of the peat soil is on average lower than in a wet year by 6-10 ° C. In the abnormally hot summer (2012), only the upper 20 cm of soil was warmer than in the over moistened year (2018). The rest of the peat deposit was colder in 2012.

- Drainage of a bog leads to a decrease in the levels of bog waters, and as a result, a decrease in soil temperature in both summer and winter.



# THX FOR ATTENTION

