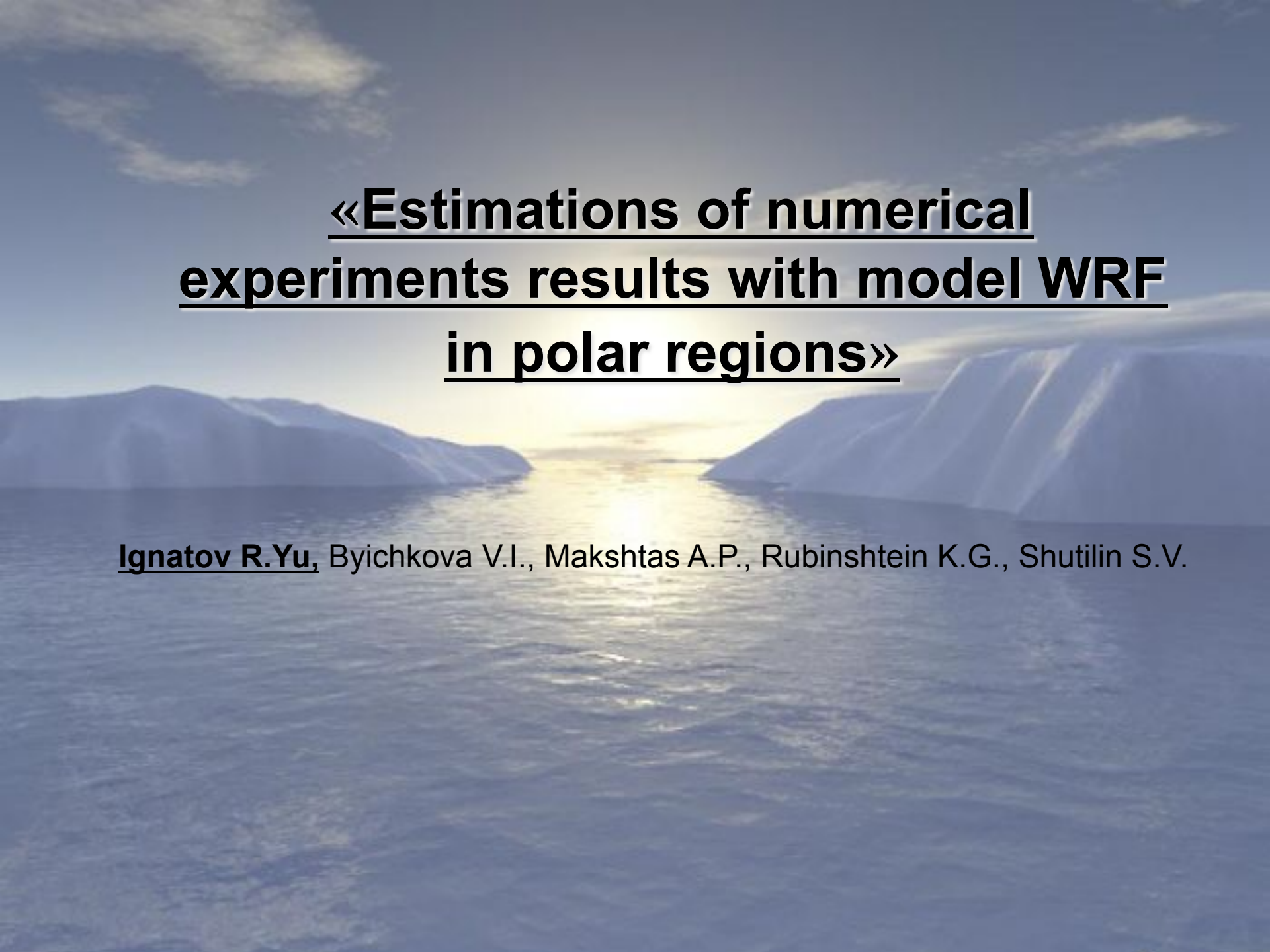




«Estimations of numerical experiments results with model WRF in polar regions»

Ignatov R.Yu, Byichkova V.I., Makshtas A.P., Rubinshtein K.G., Shutilin S.V.

Non-hydrostatic regional model WRF

(The national centre of Atmospheric Researches),

Classic

Polar

Features of the polar version

Albedo

Density

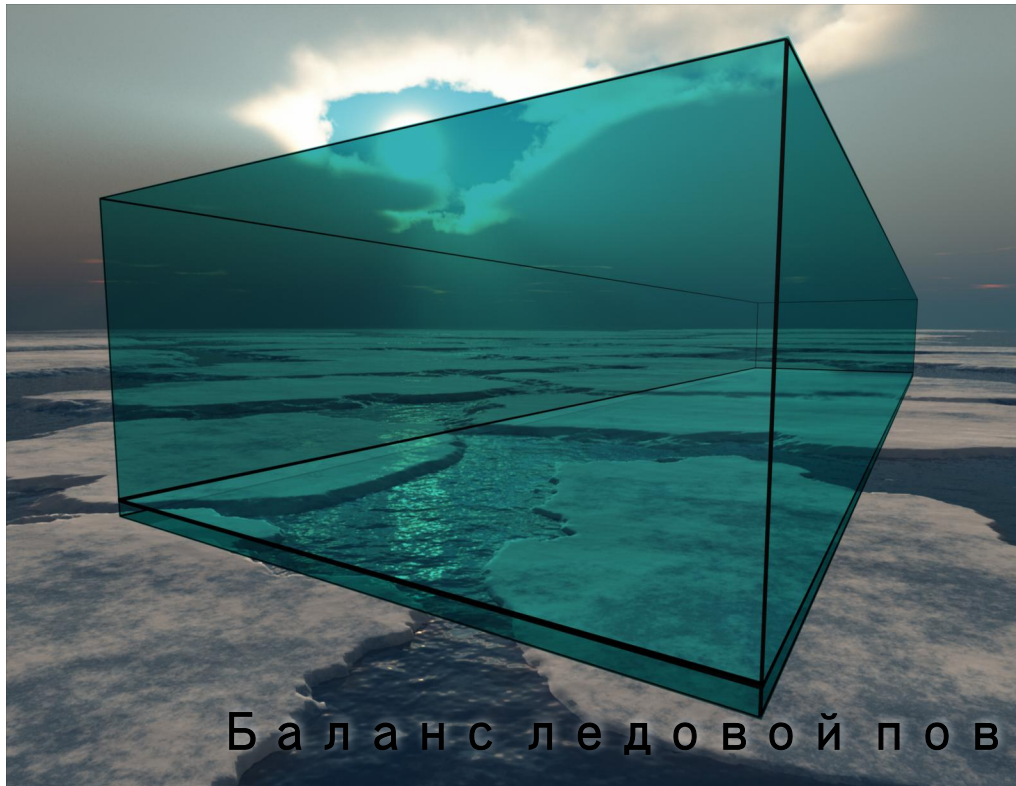
Heat conduction

Ice fraction

Thermometric
conductivity

Water content

Roughness

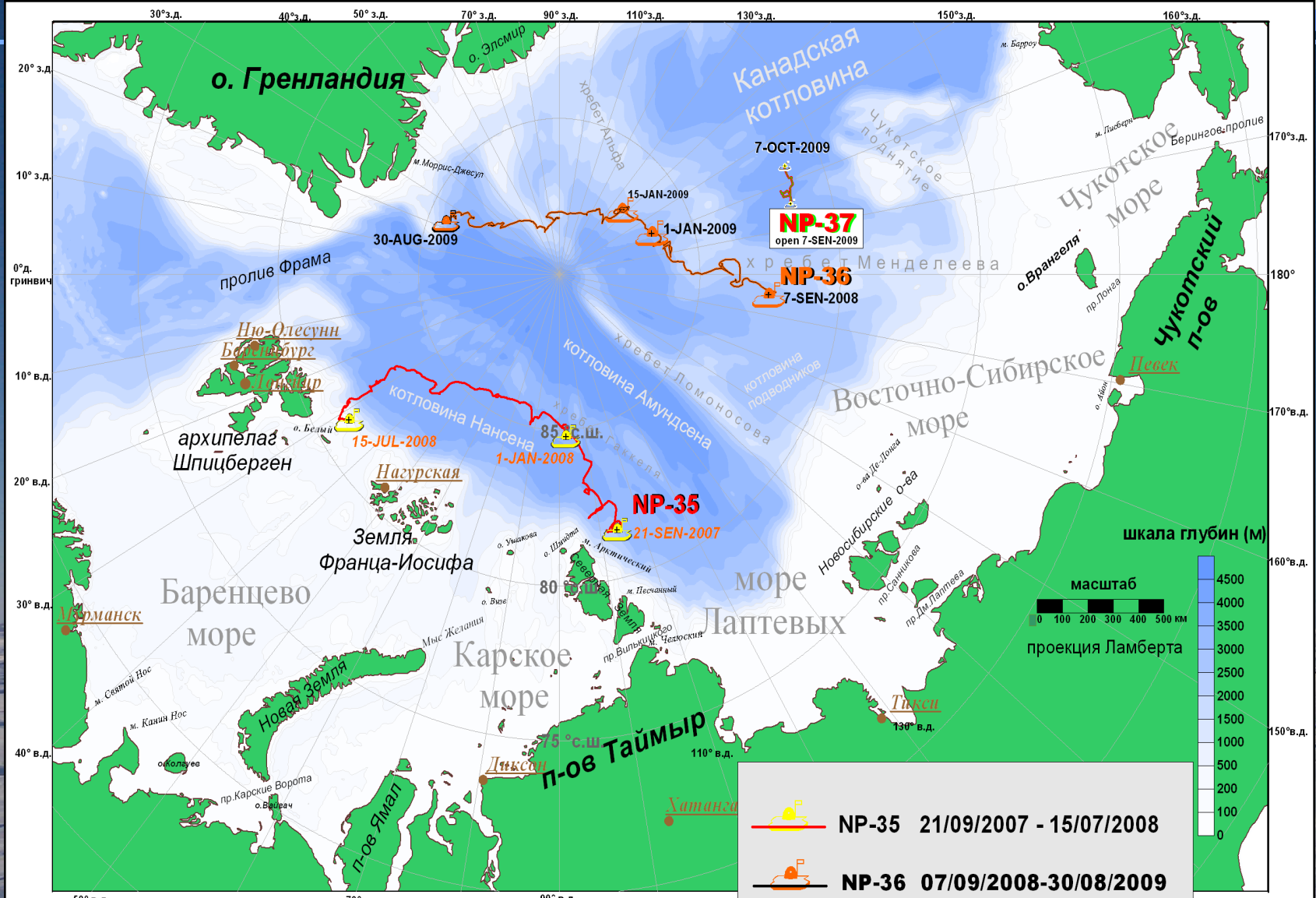


Б а л а н с л е д о в о й п о в е р х н о с т и

Topic

Verification of model WRF on the basis of unique observations given from drifting stations the North Pole – 35 and the North Pole – 36.

An estimation sensible and latent fluxes from underlying surface.

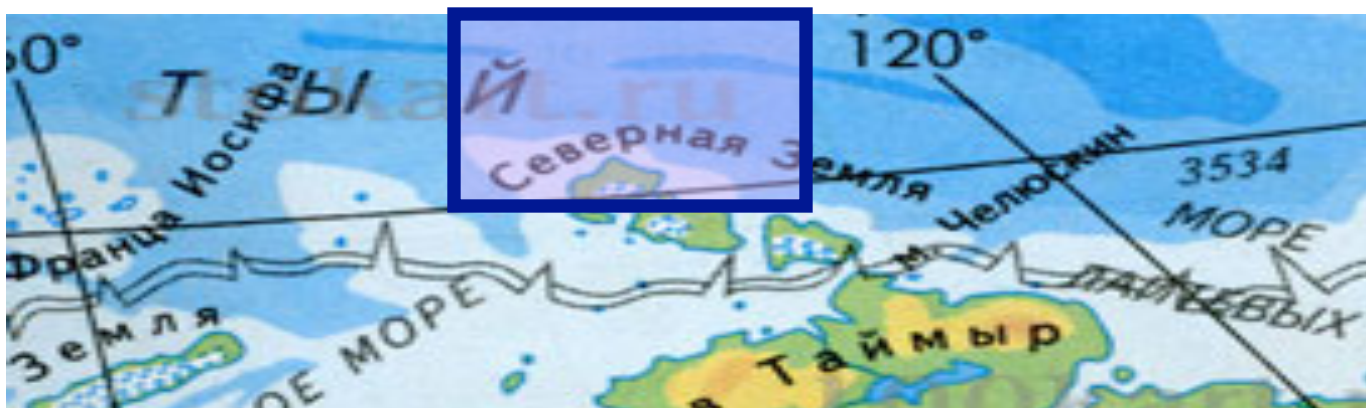


HIGH LATITUDE ARCTIC EXPEDITON OF AARI, ROSHYDROMET



Preliminary estimations of model WRF

11.12.2007 - 19.12.2007



Spatial step - 2 км,
Length of forecast - 36 h.

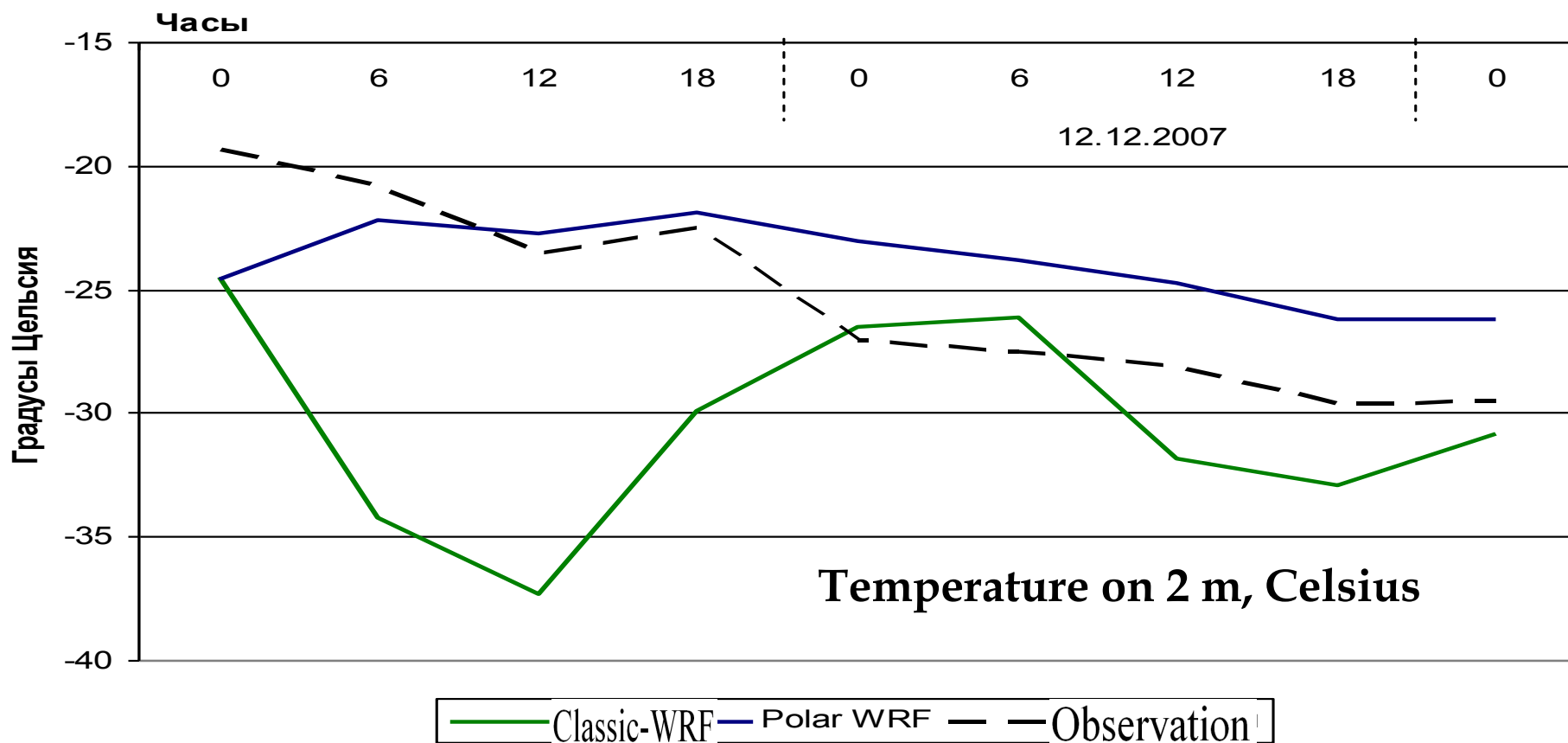
Initial and boundary conditions - NCAR Finally Analysis

Estimation of initial fields, mean for 11.12.2007 - 19.12.2007

	Mean difference	Standard deviation
Temperature on 2м (C)	4.7	2.9
Humidity on 2м (%)	8	3
Wind speed on 10 м (м/с)	2.4	3.1

Estimations of forecasts, Polar and Classic versions of WRF,
mean for 11.12.2007 - 19.12.2007.

	<u>Polar version</u>		<u>Classic version</u>	
	Mean difference	Standard deviation	Mean difference	Standard deviation
Temperature on 2м (C)	2.7	1.4	6.8	5.3
Humidity on 2м (%)	4	2	9	6
Wind speed on 10 м (м/с)	1.4	0.6	1.4	0.6



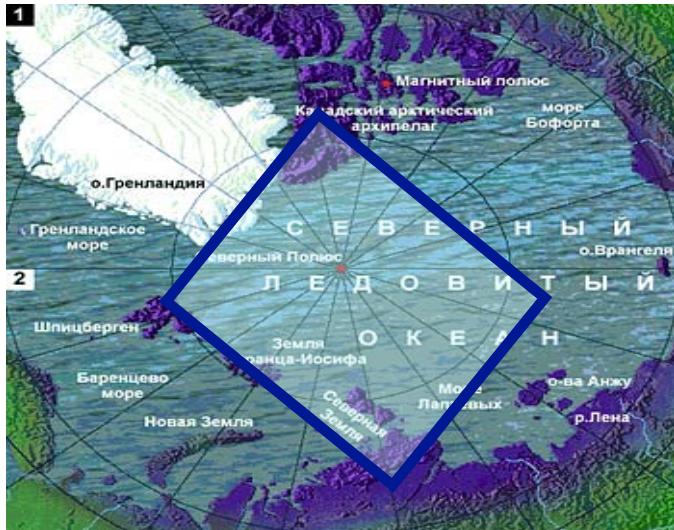
Mean estimations on a column of atmosphere from a surface to 100 hPa vertical profiles of temperature, relative humidity and wind speed for a cloudless interval and cloudy.

	Cloudlessly interval		Cloudly interval	
	Mean difference	Standard deviatiiin	Mean difference	Standard deviatiiin
Temperature on 2М (C)	2.1	1.7	1.9	1.6
Humidity on 2М (%)	8.5	7.3	10.9	11.3
Wind speed on 10 М (m/c)	1.8	1.3	1.7	1.4

Conclusions:

- Preliminary level of estimations which can improve further is received
- Calculations on the polar version of model WRF are closer to measurements than on the common version
- Initial fields produce the big Errors which influence on forecast quality
- More exact description of balance of a surface attracts considerable improvement quality of forecasts.

Выбор оптимальной конфигурации модели для полярных областей.



Parameterization of radiation

R1 - Dudhia J scheme

R2 – RRTMG Scheme

Parameterization of turbulence

T 1 - Monin-Obufhov scheme

T 2 – Monin-Obukhov (Janjic) scheme

Parameterization of boundary layer

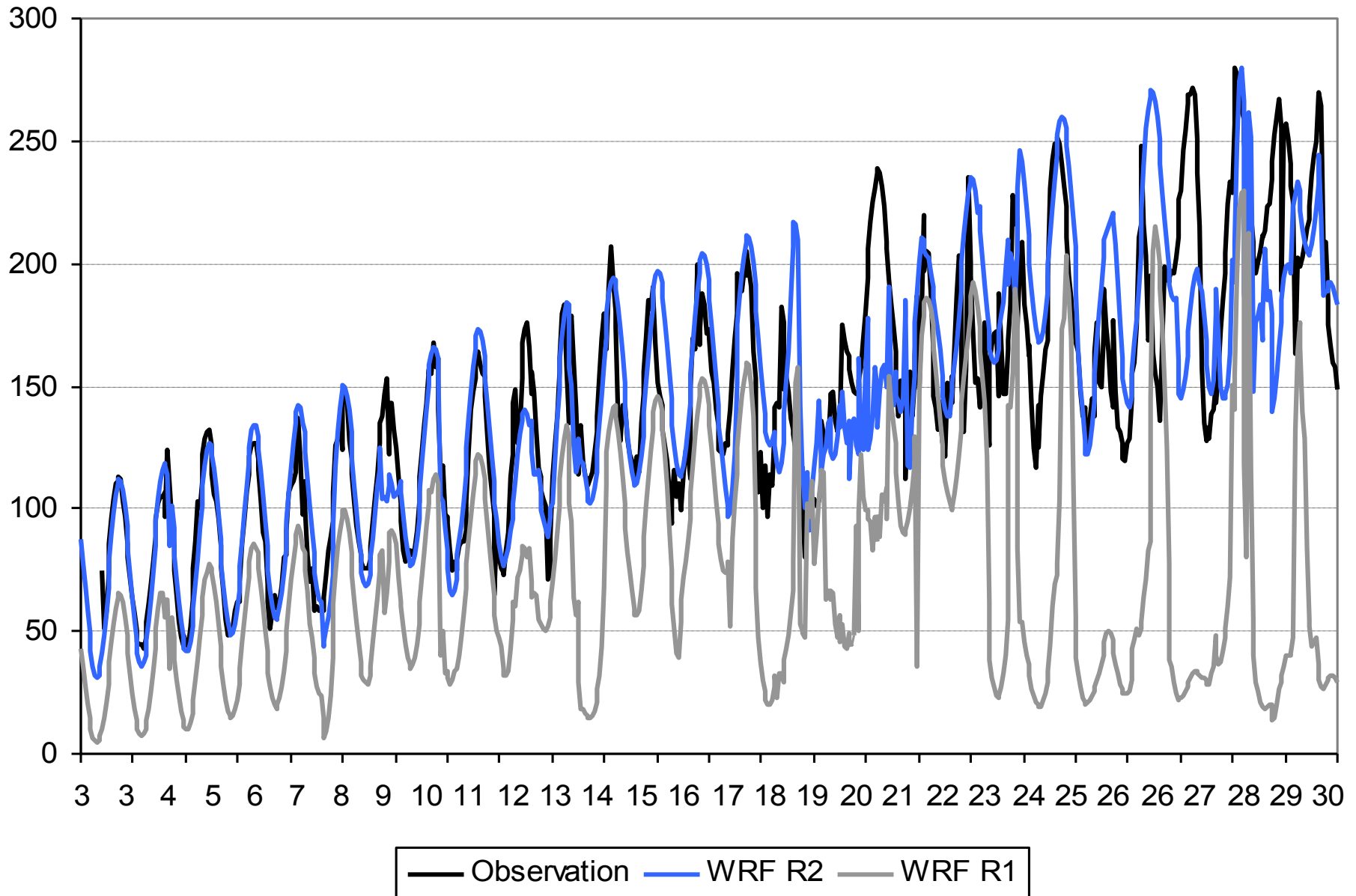
P1 – Mellor-Yamada-Janjic TKE scheme

P2 - Mellor-Yamada-Janjic 3rd level TKE scheme

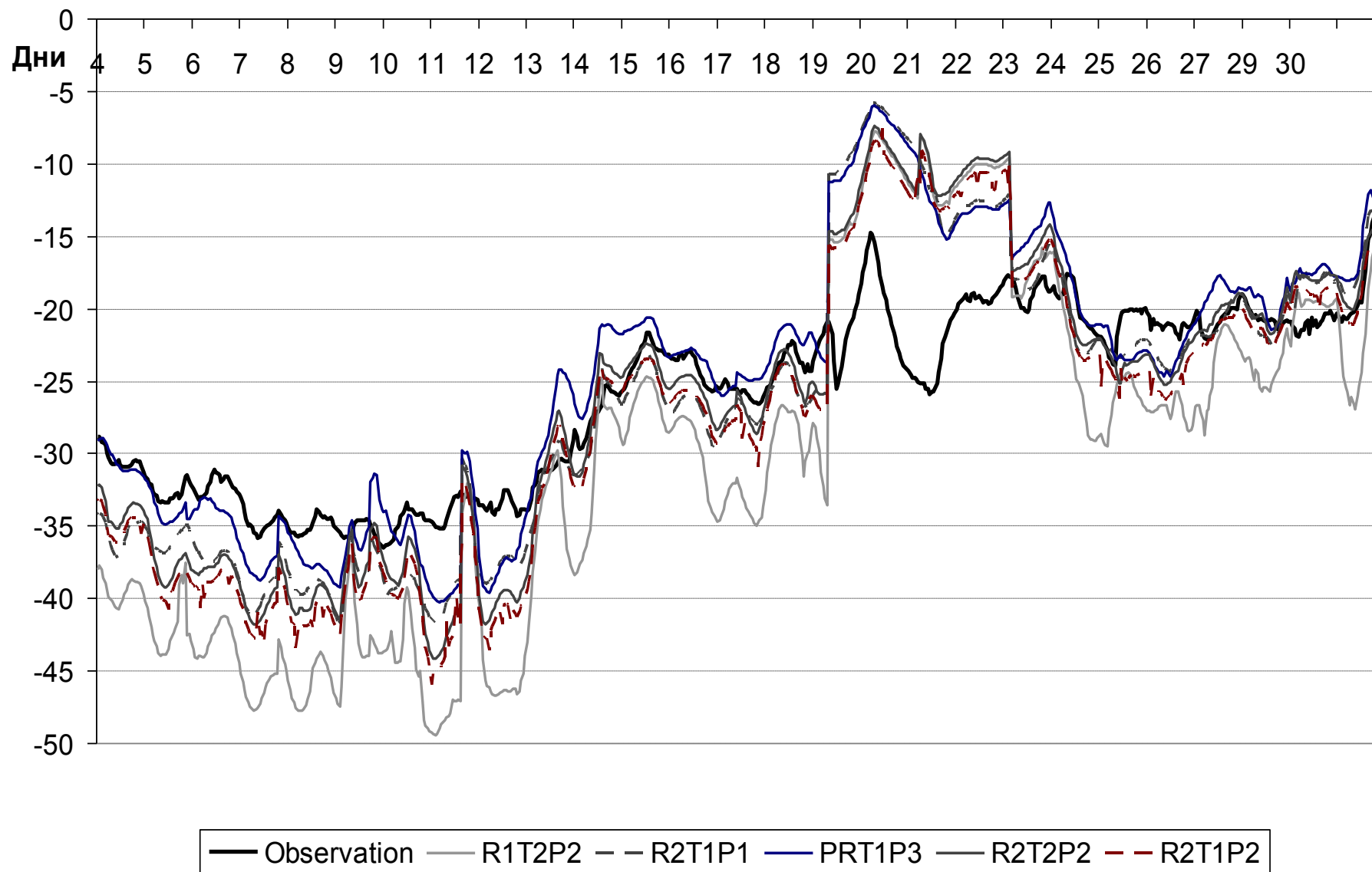
P3 – Modified Bougeault and Lacarrere scheme

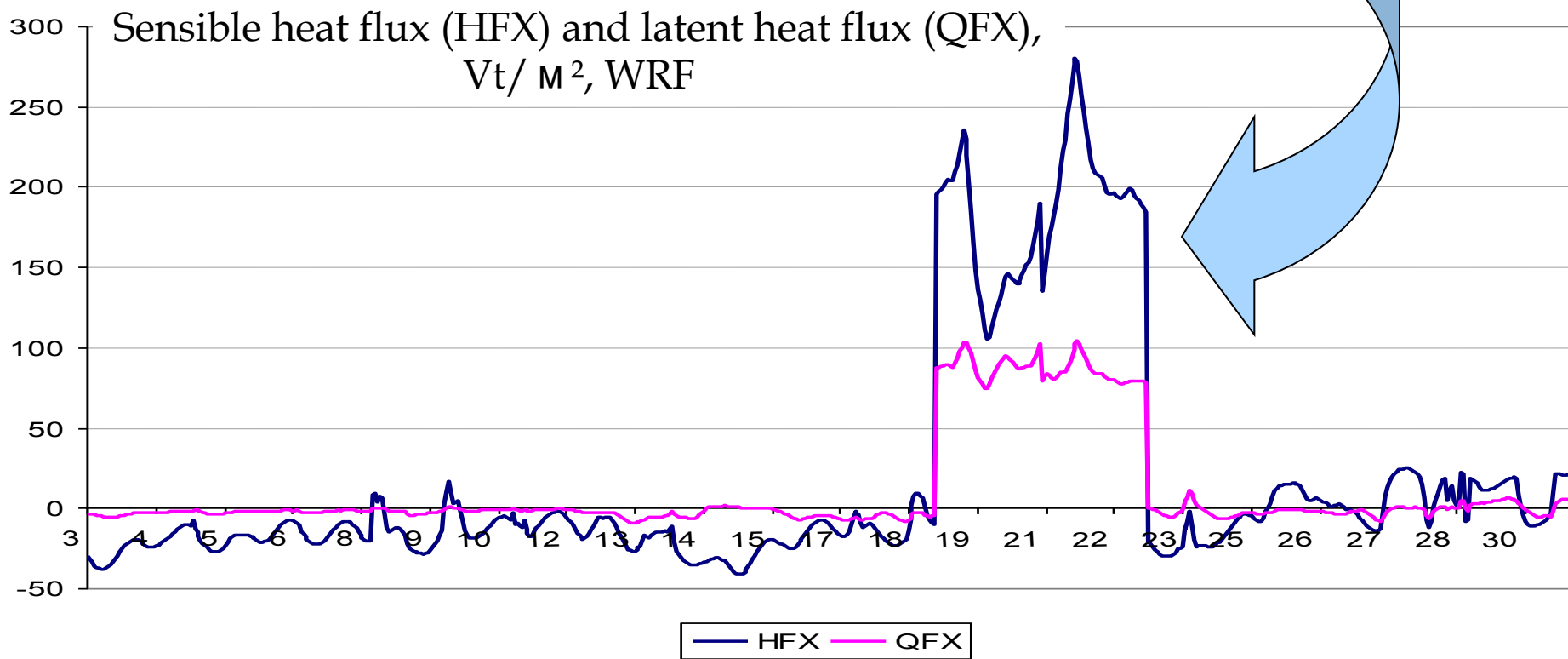
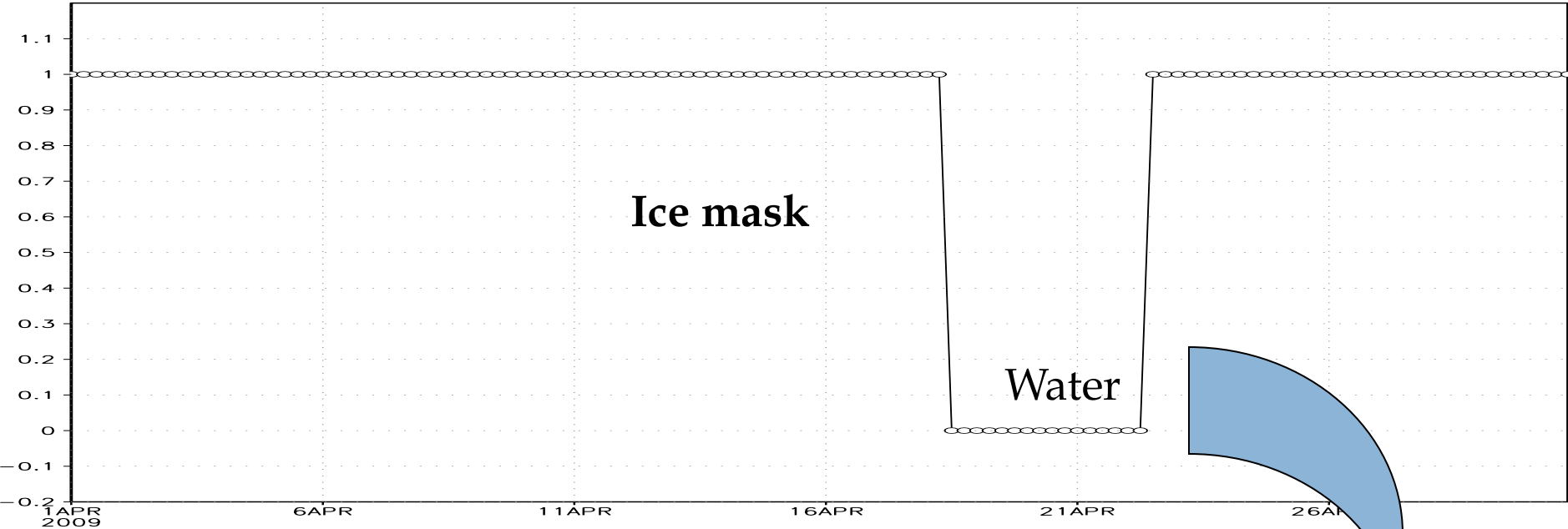
[illegible]

Coming short-wave radiation (Wt/m^2). Observation and experiments with WRF with two various parameterizations of radiation.



Temperature on 2 m (°C). Observation and experiments with different parameterizations, April 2009.





Estimations of forecast

spatial step 10 km

Experiment	T2 (°C)		T8 (°C)		W10 (m/c)	
	Δ	D	Δ	D	Δ	D
R1 T 2P2	5.4	7.8	6.3	7.6	1.2	1.7
R2 T 1P1	<u>2.6</u>	2.9	2.4	3.2	1.2	1.4
R2 T 1P3	<u>2.1</u>	2.4	2.2	2.4	2.2	1.9
R2 T 2P2	3.0	3.2	3.4	3.4	1.3	2.0
R2 T 1P2	2.9	3.3	2.5	3.5	1.7	2.0

spatial step 2 km

Experiment	T2 (°C)		T8 (°C)		Q2 (g/kg)		Q8 (q/kg)		W10 (m/c)	
	Δ	D	Δ	D	Δ	D	Δ	D	Δ	D
R2 T 1P3	<u>1.8</u>	<u>2.2</u>	1.8	2.3	0.1	0.02	0.1	0.02	1.3	1.9

Calculation of turbulence fluxes

$$Ri = 0$$

Indifferent stratification

$$\Delta U = \frac{U_*}{\chi} \ln \frac{z_2}{z_1}$$

$$\Delta \theta = \theta_* \ln \frac{z_2}{z_1}$$

$$\Delta q = q_* \ln \frac{z_2}{z_1}$$

$$Ri \geq 0.2$$

Strongly unstable stratification

$$U_* = \frac{\chi \Delta U + \chi N(z_2 - z_1)}{(\ln \frac{z_2}{z_1} + \beta_u \frac{z_2 - z_1}{L})}$$

$$\theta_* = \frac{\chi \Delta \theta}{(\ln \frac{z_2}{z_1} + C_o \frac{z_2 - z_1}{L} + C_o C_N \frac{Fr^3 (z_2 - z_1)}{L})}$$

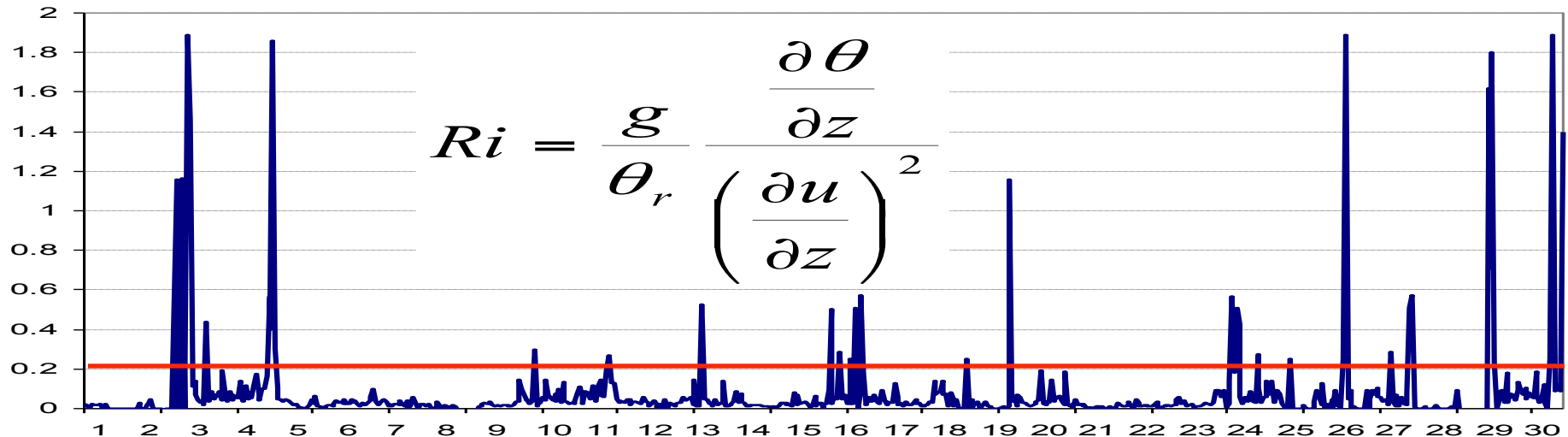
$$0.2 > Ri > 0$$

Stable stratification

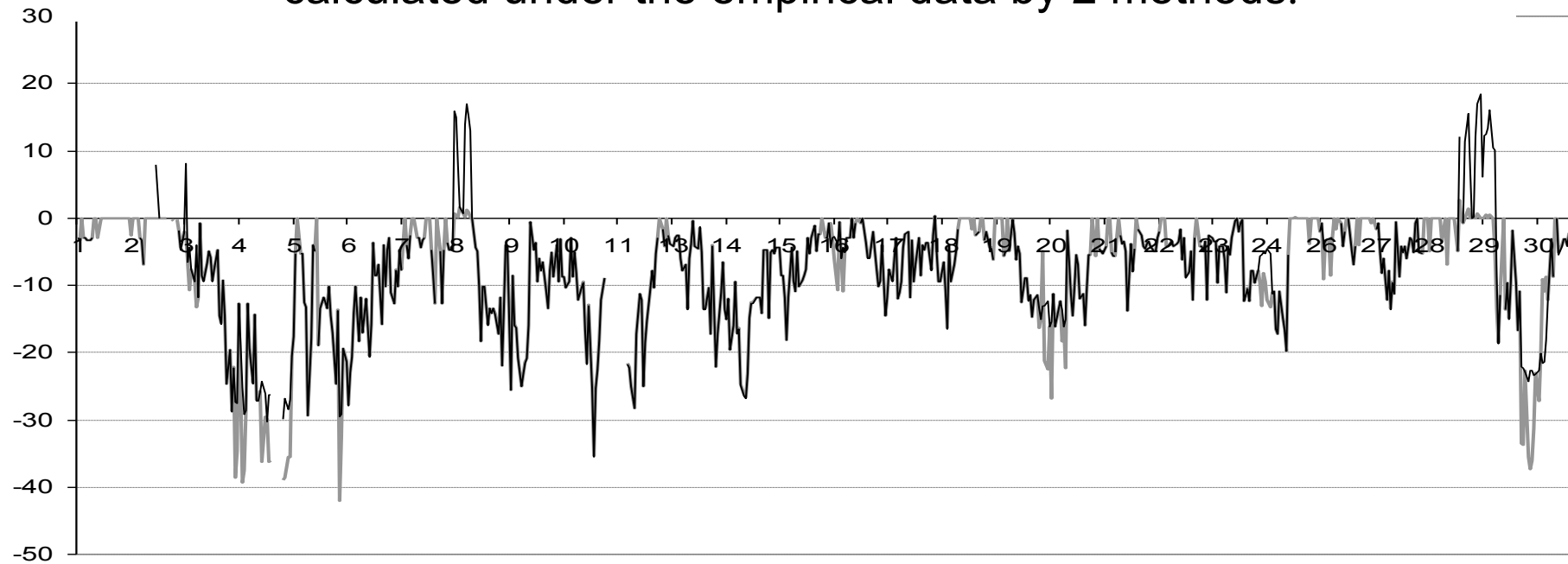
$$\Delta U = \frac{U_*}{\chi} (\ln \frac{z_2}{z_1} + \beta_u \frac{z_2 - z_1}{L})$$

$$\Delta \theta = \theta_* (\ln \frac{z_2}{z_1} + \beta_\theta \frac{z_2 - z_1}{L})$$

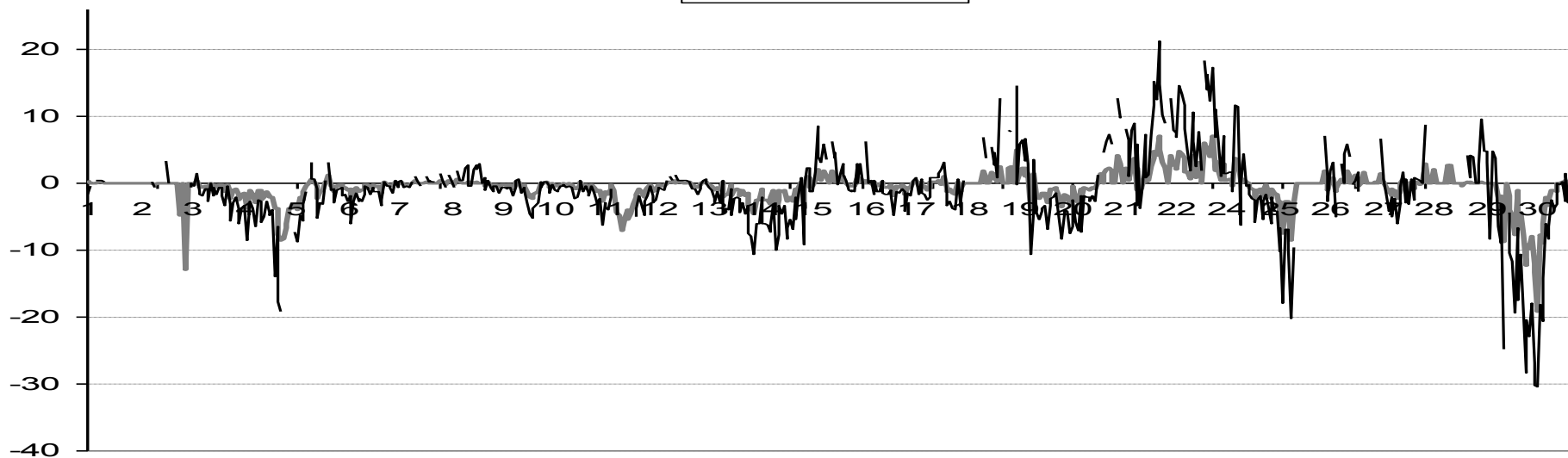
$$\Delta q = q_* (\ln \frac{z_2}{z_1} + \beta_\theta \frac{z_2 - z_1}{L})$$



Turbulent fluxes of the sensible (HFX) and latent (QFX) heat, calculated under the empirical data by 2 methods.

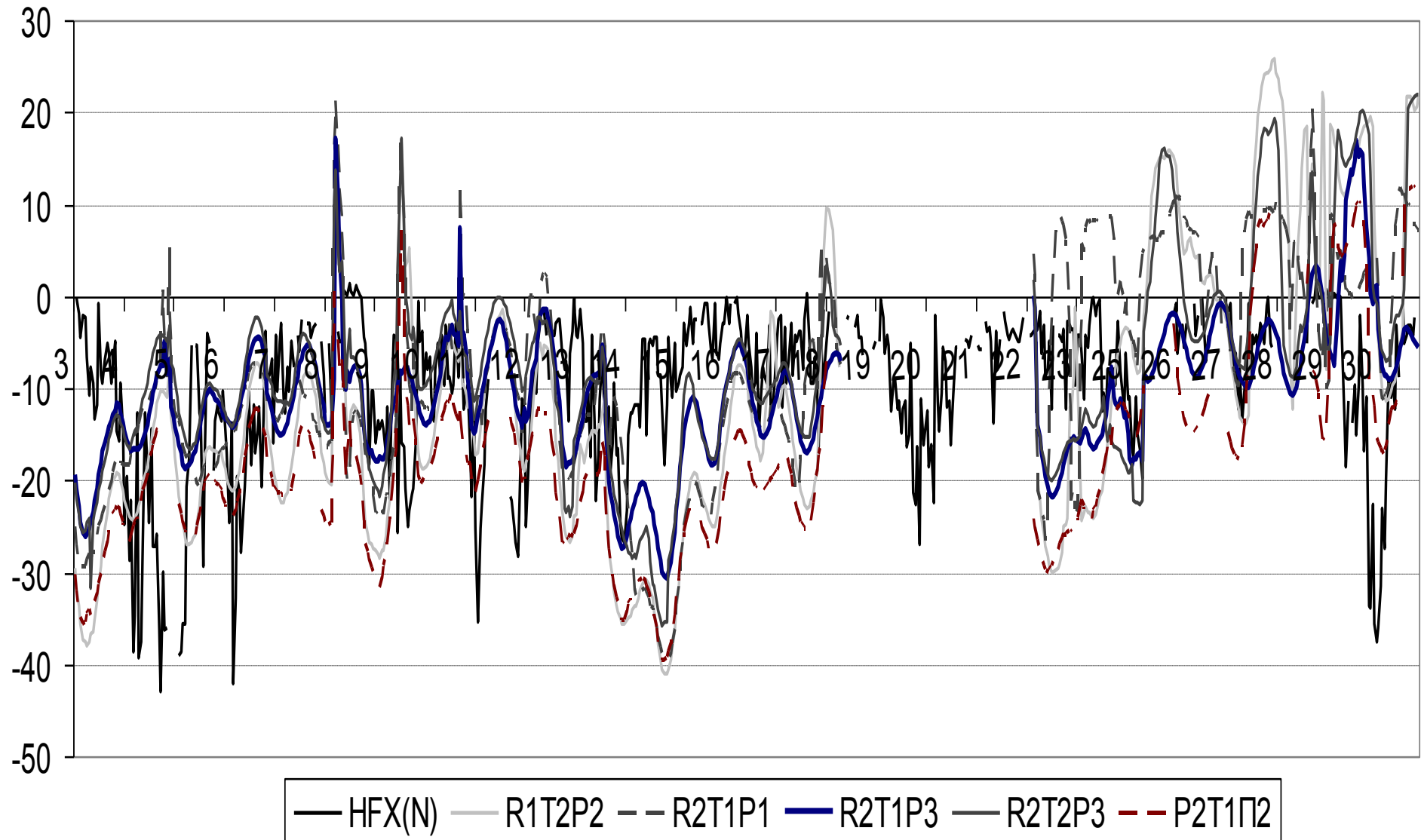


— HFX(N) — HFX



— QFX(N) — QFX

Sensible heat fluxes (Wt/m^2), calculated by observation and experiments with model WRF with various parameterization.



Errors of turbulent flux of sensible and latent heat flux
(experiments with WRF-model with different parameterization)

spatial step 10 km

Experiment	HFX (BT/M ²)		QFX (BT/M ²)	
	Δ	D	Δ	D
R1 T 2P2	13,2	80,0	2,8	7,3
R2 T 1P1	<u>9,0</u>	<u>57,0</u>	2,6	8,6
R2 T 1P3	<u>8,4</u>	<u>48,0</u>	2,8	11,1
R2 T 2P2	13,0	65,0	2,7	11,1
R2 T 1P2	11,6	61,0	3,1	8,2
HirHam	14,6	110,0	4,2	19,0

spatial step 2 km

Experiment	HFX (BT/M ²)		QFX (BT/M ²)	
	Δ	D	Δ	D
R2 T 1P3	7.6	<u>36</u>	2.7	9.7

Conclusions

- Validation of regional mesoscale model WRF by the drifting stations "North Poles" data was done for the first time. Base level of estimations of forecasts of the basic meteorological values in Arctic regions is received. It is shown that for forecasting of weather in Arctic regions it is closer to measurements using the polar version of model WRF.
- The optimum configuration of the model WRF for polar regions is found.
- Two methods of turbulent fluxes calculation under the empirical data are analysed. Monin-Obukhov method of calculation including components containing information on stratification of atmosphere is accepted.
- On the basis of the calculated values under the empirical data of fluxes heat and latent heat in models WRF and HIRHAM have been estimated. It was that the forecast quality of turbulent fluxes depends on stratification.



**Thank you
for attention!**

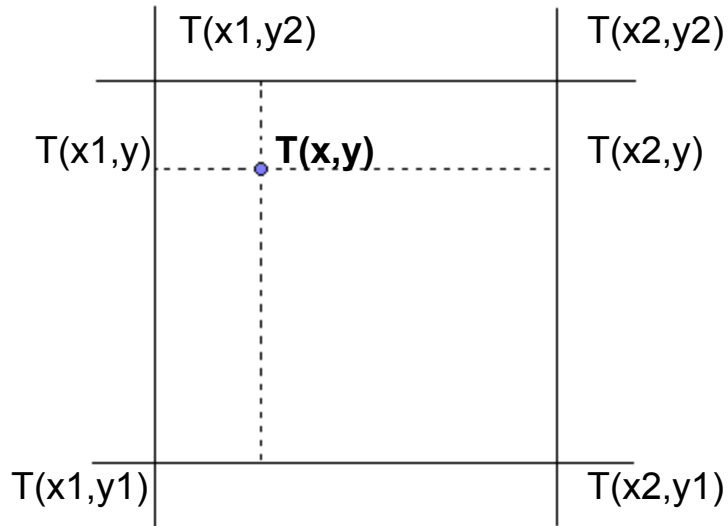
$$H = -\theta_* \chi c_P \rho U_*$$

$$LE = -\rho L q_* U_*$$

$$N^2 = \left(B \frac{\partial \theta}{\partial z} + 0.61 g \frac{\partial \theta}{\partial z} \right)$$

$$B = \frac{g}{T}$$

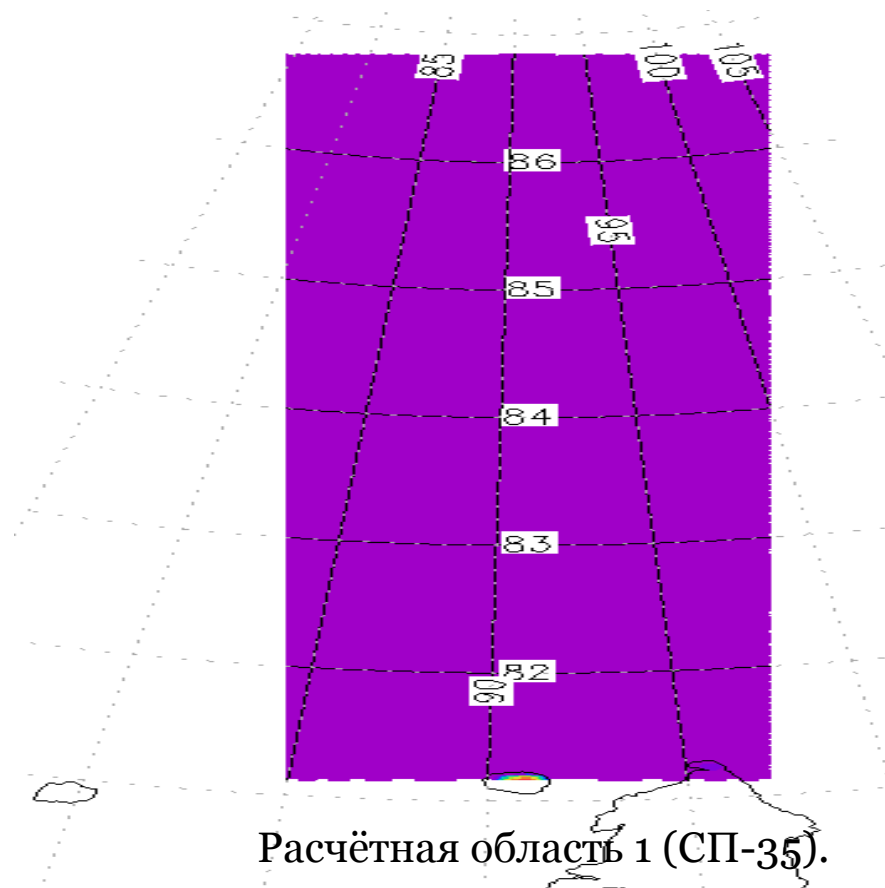
$$Fr = \frac{LN}{U^*}$$



$$T(x1, y) = T(x1, y2) \left(\frac{y - y1}{y2 - y1} \right) + T(x1, y1) \left(\frac{y2 - y}{y2 - y1} \right)$$

$$T(x2, y) = T(x2, y2) \left(\frac{y - y1}{y2 - y1} \right) + T(x2, y1) \left(\frac{y2 - y}{y2 - y1} \right)$$

$$T(x, y) = T(x1, y) \left(\frac{x2 - x}{x2 - x1} \right) + T(x2, y) \left(\frac{x - x1}{x2 - x1} \right)$$



облачно –с 19.01.2008 по 24.01.2008

безоблачно – с 11.12.2007 по 19.12.2007

Ri (0.2), разделение стратификации по числу

Skamarock W. C. et al. A description of the Advanced Research WRF version 2.
NCAR Tech. Note No.468, June 2005 (last revision January 2007), 88 pp.

Расчёт турбулентных потоков.

Zilitinkevich, S. S., Perov, V. L., and King, J. C.: Near-surface turbulent fluxes in stable stratification: calculation techniques for use in general circulation models, Quart. J. Roy. Met. Soc., 128, 1571–1587, 2002.