

Long-term changes of stationary waves in the atmosphere

Oznobikhina O.A., Mordvinov V.I., Devyatova E.V.,
Institute of Solar-Terrestrial Physics, SB, RAS, Irkutsk
ozn-olesya@yandex.ru

The goals of this work were:

1. Analysis of the spatial structure and long-term changes in stationary planetary waves in the Northern Hemisphere.
 2. Calculation of the structure of stationary waves due to orographic forcing, using quasi-geostrophic barotropic model.
- Comparison of modeling results with observational data.

Figures show the distributions of the 500 hPa height in July, averaged over the decades 1950-1960 (Fig. 1) and 2000 - 2010 (Fig. 2).

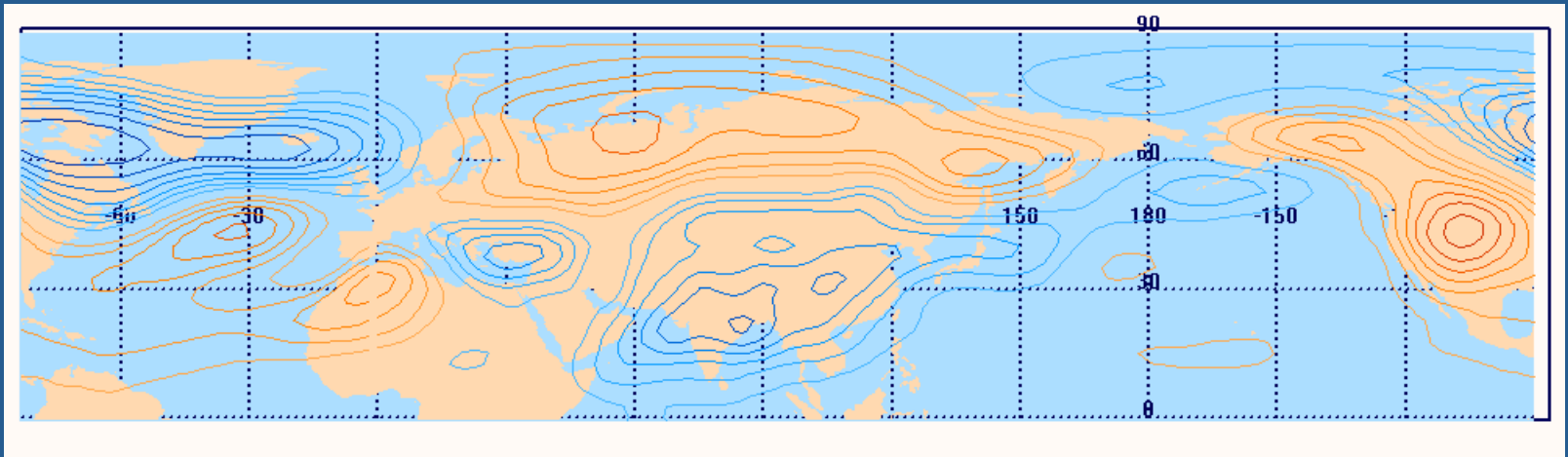


Fig. 1 – Planetary waves at 500 gPa in July 1950-1960. The distance between contours of 10 m

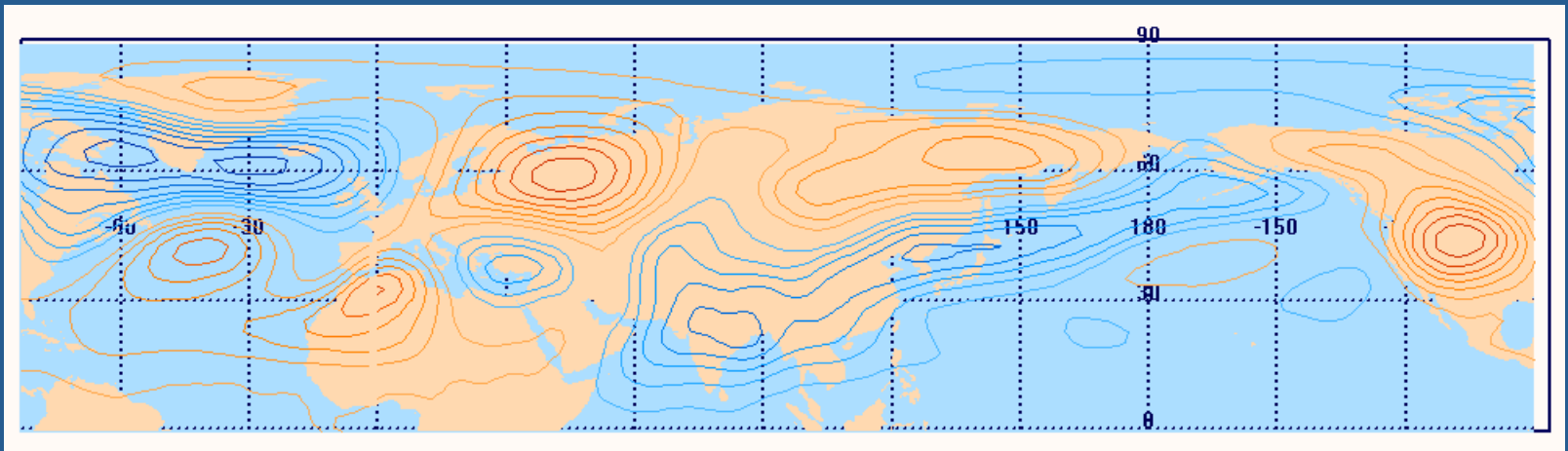


Fig. 2 – Planetary waves at 500 gPa in July 2000-2010.

Asia - Europe - Atlantic

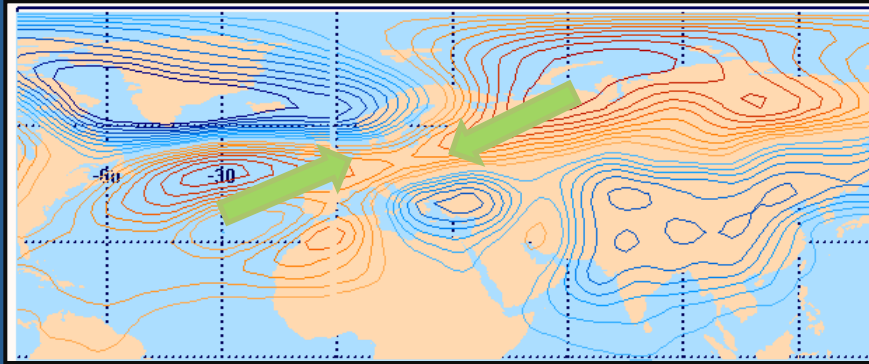


Fig. 3 – The observed structure of planetary stationary waves at 500 hPa in July 1964

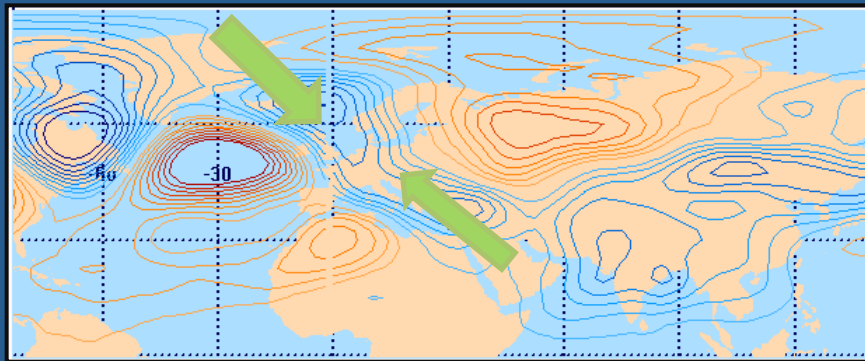


Fig. 4 – The observed structure of planetary stationary waves at 500 hPa in July 1966

Asia - Pacific - North America

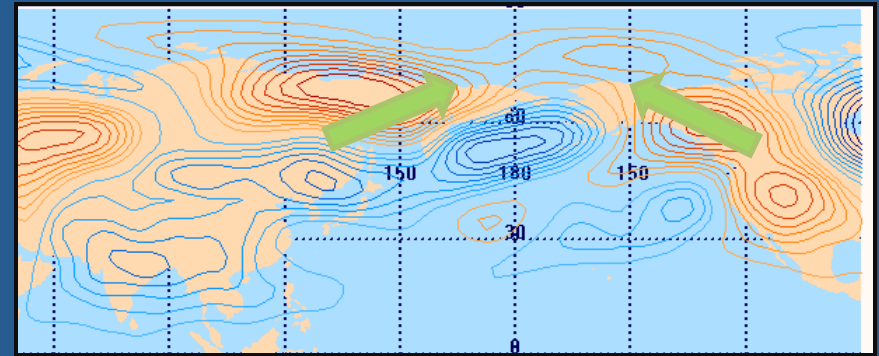


Fig.5 - The observed structure of planetary stationary waves at 500 hPa in July 1971

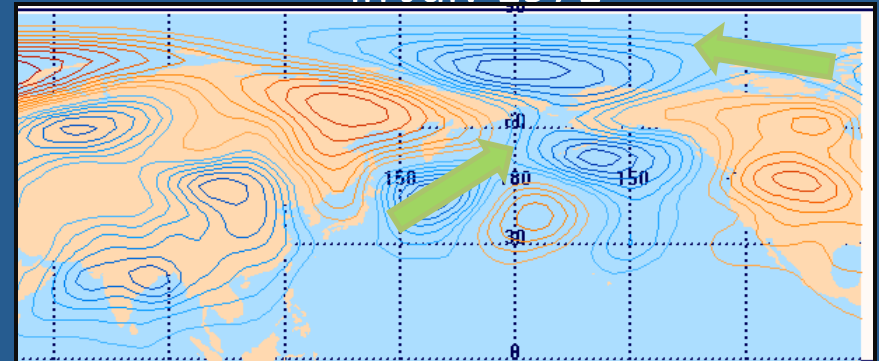


Fig.6 - The observed structure of planetary stationary waves at 500 hPa in July 1973

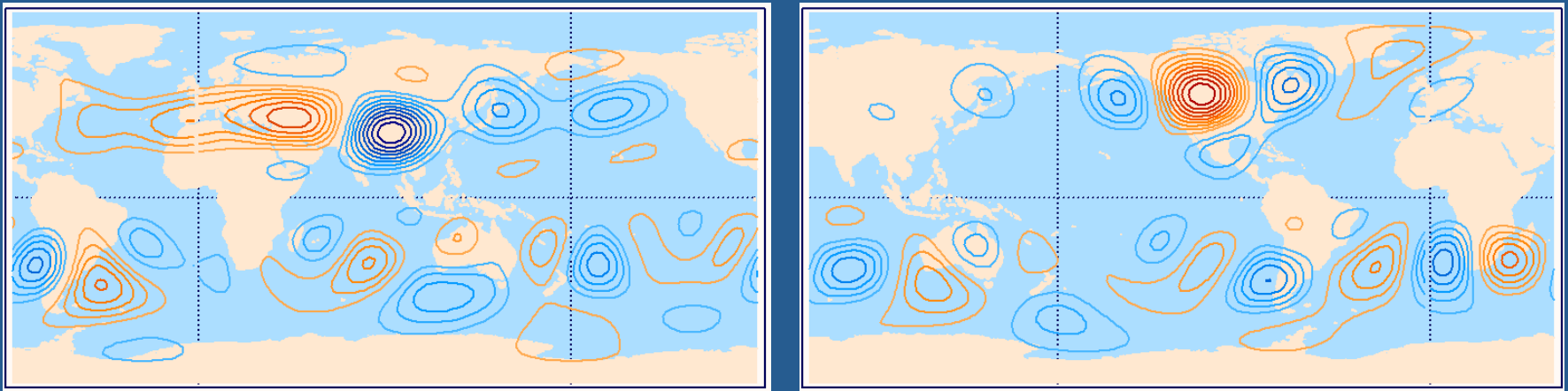


Fig.7 - Simulation of orographic forcing of stationary waves in stream function by Himalayas - Tibet (left figure) and the Cordillera - Rocky Mountains (right figure) in July 1950-1960 for the velocity at 500 hPa

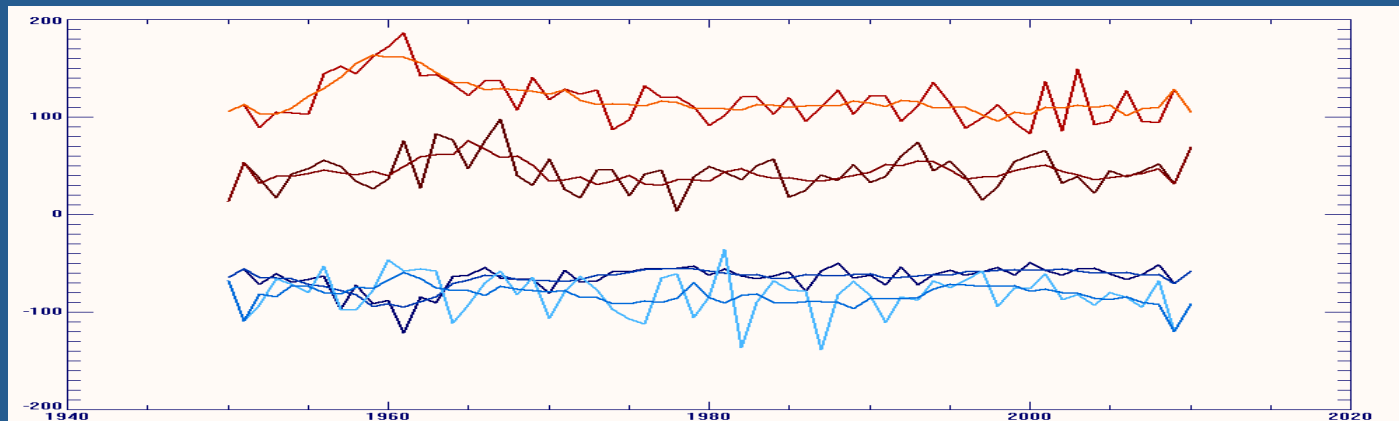


Fig. 8 - Variations of the intensities of the positive geopotential height 500 hPa anomalies over the Hawaiian Islands (red), two negative anomalies over the Asian part of Eurasia (blue) and propagating away trough over the Pacific Ocean (blue) and orographic vorticity source Himalaya - Tibet (pink) in July, 1950-2010.

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Gosabidibula G.A., Morozhina V.I., Degryatse E.V.,
Institute of Solar-Terrestrial Physics, SB, RAS, Irkutsk
gosabidibula@yandex.ru

The paper presents the results of a study of long-term changes in stationary summer planetary waves at 500 hPa in the northern hemisphere according to the NCEP/NCAR Reanalysis from 1950 to 2010 [1]. The observed structure of planetary waves changes over the northern part of Eurasia, East Asia and the Pacific Ocean. Variations of geopotential height anomalies in these regions represent decadal damped oscillations after a strong local perturbation in the late 50s - early 80s. These fluctuations can be related to the weakening in recent decades the East Pacific monsoon. In some intervals, the structure of stationary waves changed markedly in the middle latitudes, with a characteristic period of 2-4 years. As shown by model calculations in the barotropic quasi-geostrophic approximation, only the part of the waves variations can be explained by changes in orographic forcing.

The results of this work were:

- Analysis of the wavelet structure and long-term changes in stationary planetary waves in the Northern Hemisphere.
- Calculation of the structure of stationary waves due to orographic forcing, using quasi-geostrophic barotropic model.
- Comparison of modeling results with observational data.

1) Asia - Eurasia - Atlantic

Figures show the distributions of the 500 hPa height in July, averaged over the decades 1950-1960 (Fig. 1) and 2000 - 2010 (Fig. 2). Fig. 3 illustrates changes over time the intensity of the positive anomalies over the northern part of the Eurasian continent. Analysis of variance allowed the following conclusions:

- There are not significant trend in the intensity of positive geopotential height anomalies over the northern part of Eurasia in the last decades. The anomaly over Eurasia has been long-term expressed there over Asia.
- Changed the configuration of the anomalies. One possible reason for this could be spread through from the Asian part of the European continent to Western Siberia, which provided a large positive anomaly over the northern part of Eurasia into two separate anomalies.
- Fluctuations in the smoothed curves have quasi-decadal component, but the reliability of this component is low, since the amplitude of component is comparable to the standard deviation of interannual fluctuations.

2) Asia - Pacific - North America

In the region of Asia and the Pacific, there were large-scale fluctuations in the circulation, which began with the strong perturbations in the late 50s - early 60s. Later, geopotential height variations at the different regions were less agreed and resembled the relaxation process after a strong disturbance. One consequence of this process was to change the configuration of stationary waves over East Asia. Figure 4 shows graphs of changes in the intensity of positive geopotential height anomalies over the Hawaiian Islands and two negative anomalies over East Asia and the Pacific Ocean. Significant variations of the anomalies mostly completed by the 70 in year. Reducing pressure over the East Asia was accompanied by mass transfer. First in the western Pacific, and then, with some delay, - in the region of Hawaii maximum. Correlation coefficients (Table 1) between the negative anomalies over the Pacific Ocean and Asia and between the negative anomaly over the Pacific Ocean and positive anomalies over Hawaii confirmed this conclusion.

2 - 4 year fluctuations of the same type as those of Europe - the North Atlantic, refers, and over the Pacific Ocean and the Bering Sea and the central regions of North America. On the right are examples of this variation in 1951 (Fig. 6) and 1973 (Fig. 7) respectively.

3) The basis of the model in the equation of quasigeostrophic potential vorticity (in the barotropic approximation)

$$\frac{D(\bar{\omega} - \frac{1}{f} \nabla^2 \bar{\psi})}{Dt} = \frac{1}{f} \frac{\partial \bar{\omega}}{\partial t} \frac{\partial \bar{\psi}}{\partial t} - \frac{\partial \bar{\omega}}{\partial t} \frac{\partial \bar{\psi}}{\partial t} + \frac{\partial \bar{\omega}}{\partial t} \frac{\partial \bar{\psi}}{\partial t} - \frac{\partial \bar{\omega}}{\partial t} \frac{\partial \bar{\psi}}{\partial t}$$

where $\bar{\omega}$ - stream function, $\bar{\psi}$ - vorticity source, $f = 2\Omega \sin \varphi$, Ω - angular velocity of Earth, φ - the polar angle

Orographic source of vorticity determines the spatial structure of the geopotential height anomalies near the mountains, the intensity of which has considerably changed when the flow is located on the mountain. Fig. 9 shows examples of calculation of orographic stationary waves from different sources, 1950-2010.

Fig. 10 shows the variations of the vorticity source (West - Himalayas) and the intensity of the negative geopotential height anomalies over the Asian part of Eurasia and the Pacific Ocean and positive anomalies over the Hawaiian Islands. It is seen that the variation of the vorticity source and negative anomaly are closely related (correlation coefficients for unfiltered data - 0.39 and for smoothed data - 0.37). The correspondence of orographic forcing had to be coherent coefficient of other geopotential height anomalies over northern Eurasia. However, the combination of changes in the intensity of these anomalies did not confirm this assumption. To a lesser extent on the variation of static vorticity waves along the Cordillera mountain range - the Rocky Mountains.

4) Conclusion

1. Kulevskiy I., Kanamitsu M., Kikler R. et al. The NCEP/NCAR 40-year reanalysis project. - Bull. Amer. Meteor. Soc., 1996, V. 77, pp. 427-471.

2. Branstator G.N. Horizontal Energy Propagation in a Barotropic Atmosphere with Meridional and Zonal Structures. J. Atmos. Sci., 1983, V. 40, No. 7, pp. 1688-1708.

1. Kalnay E., Kanamitsu M., Kistler R. et al. The NCEP/NCAR 40-year reanalysis project. - *Bull. Amer. Meteor. Soc.*, 1996, V. 77, pp. 437-471.

2. Stanenkov O.N. Horizontal Energy Propagation in a Barotropic Atmosphere with Meridional and Zonal Structure // *J. Atmos. Sci.* 1983, V. 40, No. 7, pp. 1689-1708.

3. ... 500 hPa anomalies over the Hawaiian Islands (red), two negative anomalies over the Asian part of Eurasia (blue) and propagating away from there over the Pacific Ocean (blue) and atmospheric velocity vector. *Horvathy - Tied* (on July 1999-10-05).