

# Parameterization of the thermal conductivity of soil in the active land layer block of the INM RAS-MSU model

# Параметризация коэффициента теплопроводности почвы в блоке деятельного слоя суши модели ИВМ РАН-МГУ



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Abstract: The active land layer block of the INM RAS-MSU model uses a parameterization of heat transfer. simplified More prospective parameterizations of the thermal conductivity coefficient were selected and implemented in the computational algorithm of the model. Switching to a new method for setting the soil thermal conductivity coefficient will significantly improve the accuracy of reproducing heat transfer in the model.

Аннотация: В блоке деятельного слоя суши модели ИВМ РАН-МГУ применяется упрощенная формула расчета коэффициента Были выбраны и теплопроводности почвы. внедрены вычислительный модели более перспективные алгоритм параметризации коэффициента теплопроводности. Переход к новому способу задания коэффициента теплопроводности почвы существенно улучшит точность воспроизведения теплопереноса в модели.

### **Motivation:**

The active land layer is a significant reservoir of heat accumulation in the Earth system [Beltrami et al., 2002]

Thermal inertia of soil at depths affect atmospheric circulation over land on time scalesfrom weeks to several months [Fatichi et al., 2020]

In the models of the Earth system, it's necessary to improve parameterization of heat and moisture transfer in the soil cover

> The model of the Earth system INM RAS-MSU uses simplified parameterization of heat transfer

### **Objective:**

Improving the description of heat transfer in soil cover in the model of the Earth system of the INM RAS-MSU

### Tasks:

- Selection of prospective parameterization of the soil thermal conductivity coefficient;
- Evaluation of the effect of their implementation in the INM RAS-MSU model;
- Collection initial soil data for modelling.

## Soil thermal conductivity models:



R. Pielke [McCumber, Pielke, 1981]

 $\lambda_T = 418.7 max(exp(-P_f - 2.7), 0.00041)$ 

- Uses function of the soil moisture potential
- Empirical relationship
- Type of soil and the amount of liquid moisture are not accounted implicitly
- Porosity, density, structural features, and mineralogical composition have only an indirect effect

### In soil physics

Johansen [Johansen, 1975]

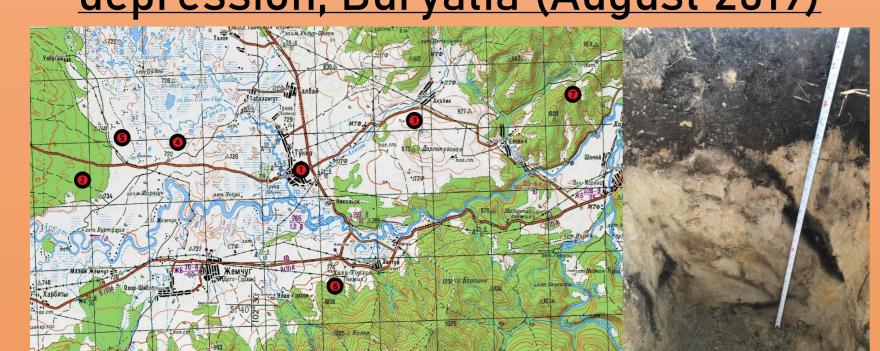
 $\lambda_{\rm T} = (\lambda_{\rm w}^{\rm n} \lambda_{\rm s}^{\rm 1-n} - \lambda_{\rm dry}) k_{\rm r} + \lambda_{\rm dry}$ 

Cote-Konrad [Cote, Konrad, 2005]

 $\lambda_{T} = \left(\lambda_{w}^{n} \lambda_{s}^{1-n} - \chi 10^{-\eta n}\right) k_{r} + \chi 10^{-\eta n}$ 

- + Based on a theoretical model
- + Parameters are available for measurement
- + Consider porosity, density
- + Developed the Johansen model
- + Parameters are available for measurement
- + Consider soil type and pore space structure

## Field studies in the Tunkinsky depression, Buryatia (August 2019)



- Collected data:
- Soil density, porosity, initial vertical moisture distribution
- **7 points** over the basin
- Depths up to 50 cm with 10 cm step
- 3 soil types: silt and clay, sandy soil, organic soil

Compared with unique measurements of soil temperature at depths up to 10 meters with 1 hour time step

## Conclusions:

- R. Pielke parameterization used in the basic version of the model less accurate than modern more parameterizations of thermal conductivity.
- The fractal model of thermal conductivity has a greater potential for improving accuracy, but it's limited by the lack of global databases for a number of its parameters.
- For the global grid of the earth system model INM RASchoosing MSU, is proposed the **Cote-Konrad** parameterization of the soil thermal conductivity.

# Thermal conductivity of porous media

Fractal model [Qin et al., 2018]

 $\lambda_{\mathrm{T}} = \gamma + \beta \frac{\gamma(\lambda_{\mathrm{s}} - \gamma)}{\lambda_{\mathrm{s}} + 2\gamma}$ 

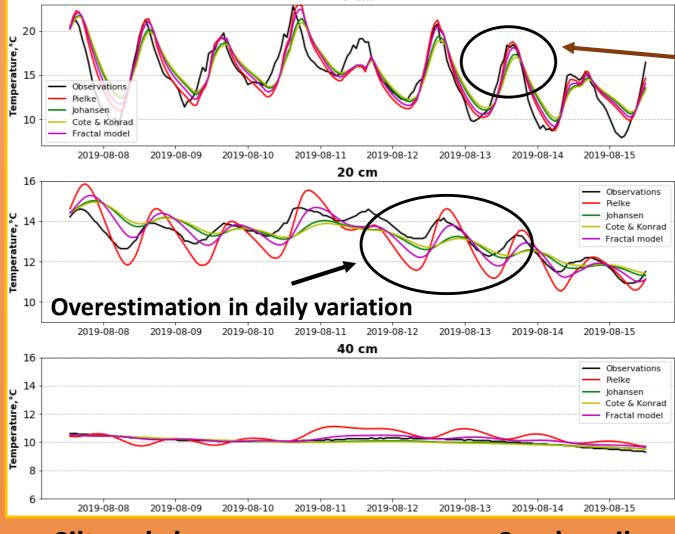
- + Based on the Fourier's law
- + Input parameters only porosity, moisture content
- + Include calibration parameters
- Restrictions for using on global grid

 $\frac{D_{Tw} + D_{fw} - 1}{D_{fw}} S_r^{1 + \frac{D_{Tw} - 1}{D_E - D_{fw}}},$   $\frac{1}{3} \frac{n}{n} \gamma = [\alpha \lambda_w + (1 - \alpha) \lambda_a]$ 

Was firstly performed for the Earth system models

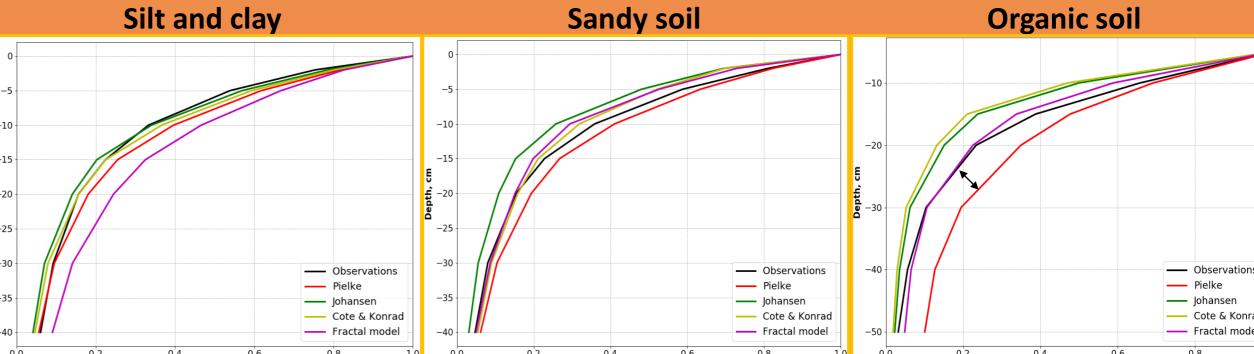
## Numerical experiments and results

Time variability in organic soil



- The surface temperature values weakly depend on the parameterization of the thermal conductivity of soil.
- Maximum errors were observed near the surface due to inaccuracies in modeling of heat balance of the surface.

Based on the Fourier's law were introduced new error metric  $\sigma_i/\sigma_{surface}$ 



For all soil types the Pielke parameterization didn't produce the most accurate results Choosing the Cote-Konrad parameterization allows to obtain the most approximate results